

Oxygen isotope dating the Australian regolith: A review and new applications

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Florian Dux

David Wheeler

Tectonic position of Australia; its translation northwards

(What is deep or intense weathering?)

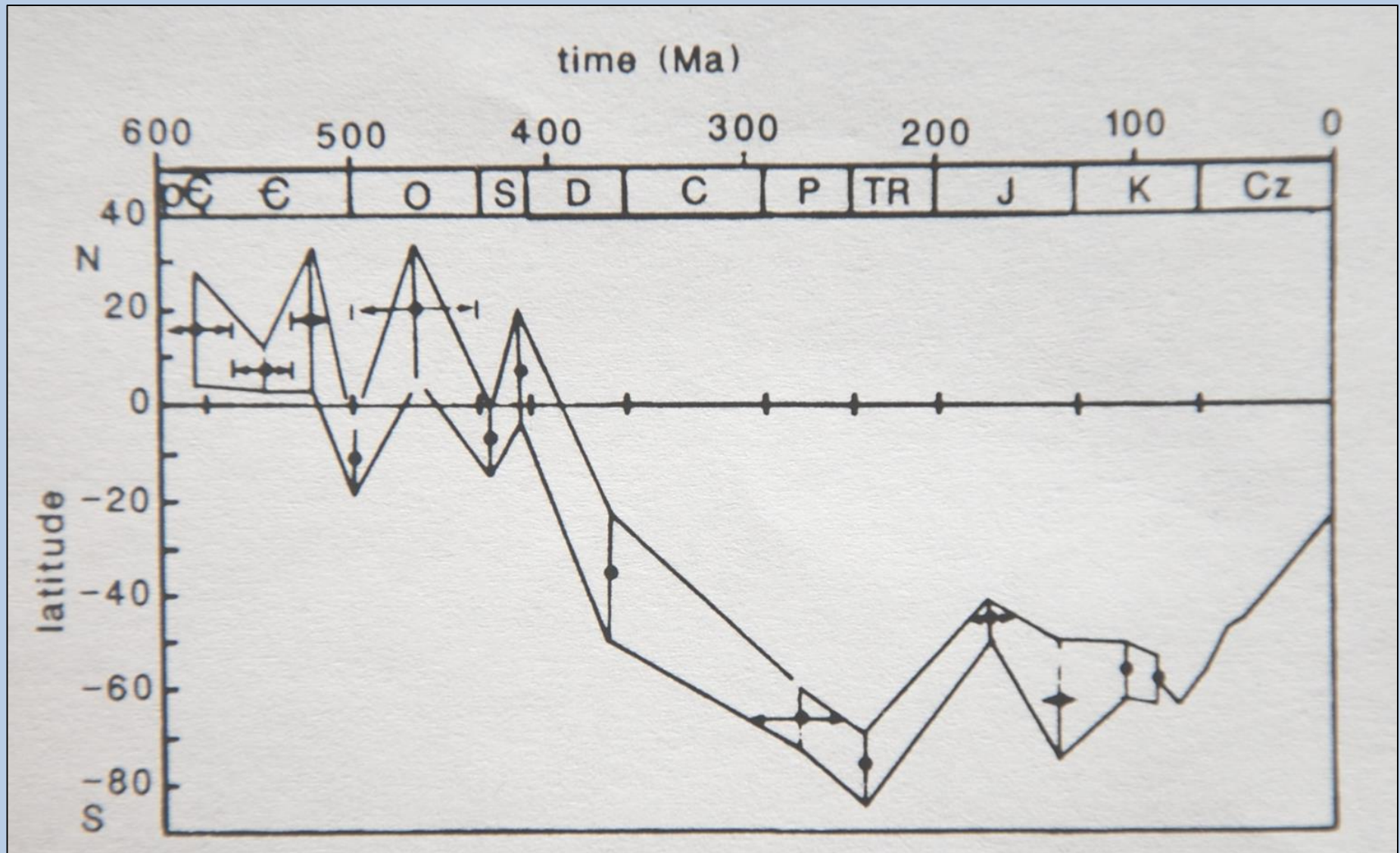
The global meteoric water line for oxygen isotopes in water

Oxygen isotopes in kaolinite of weathering origin

Ages of weathering in Australia

Examples: Yilgarn Craton and then inselbergs in the landscape

Iron oxides



(modified from Veevers, 1984)

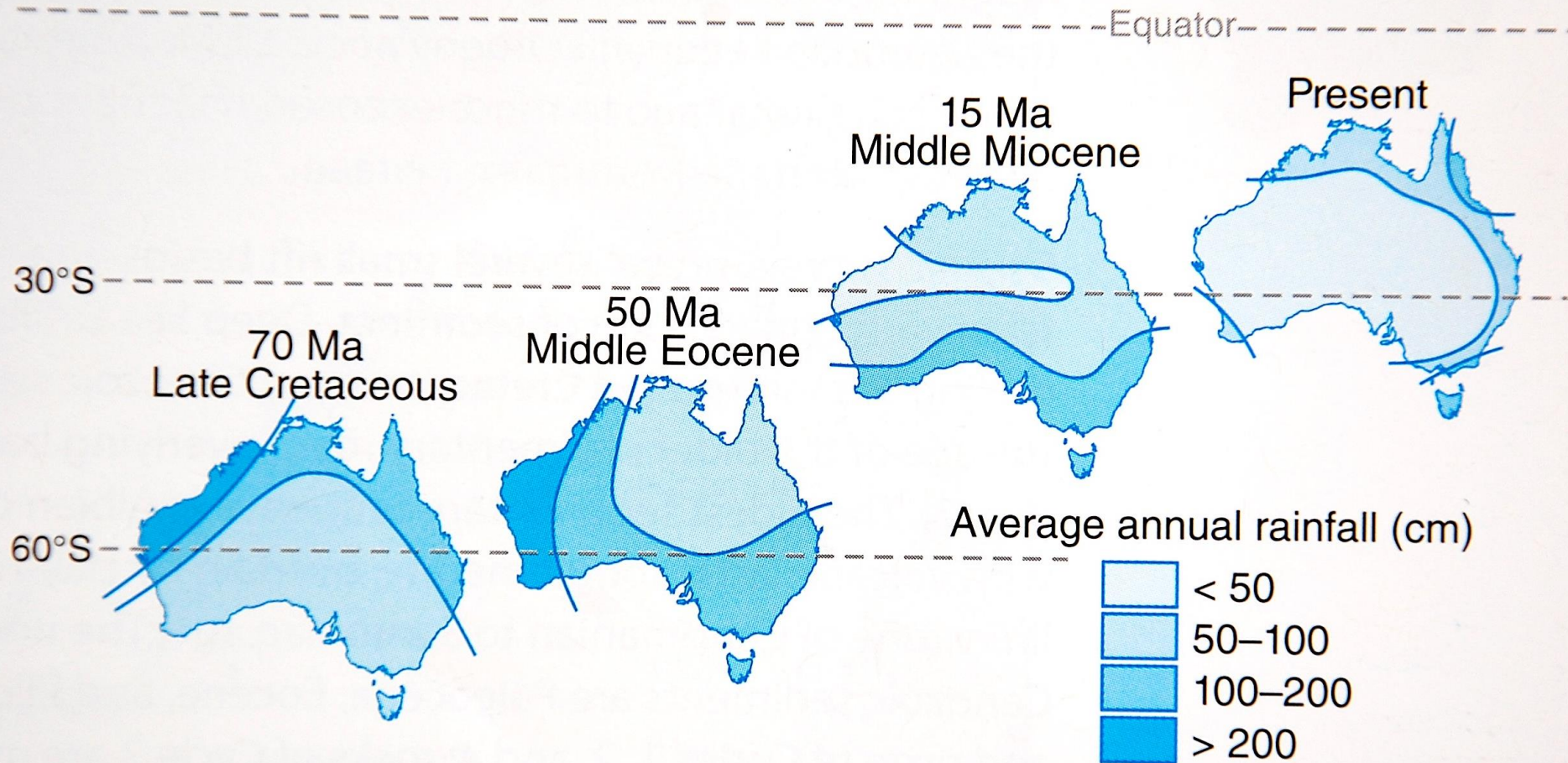
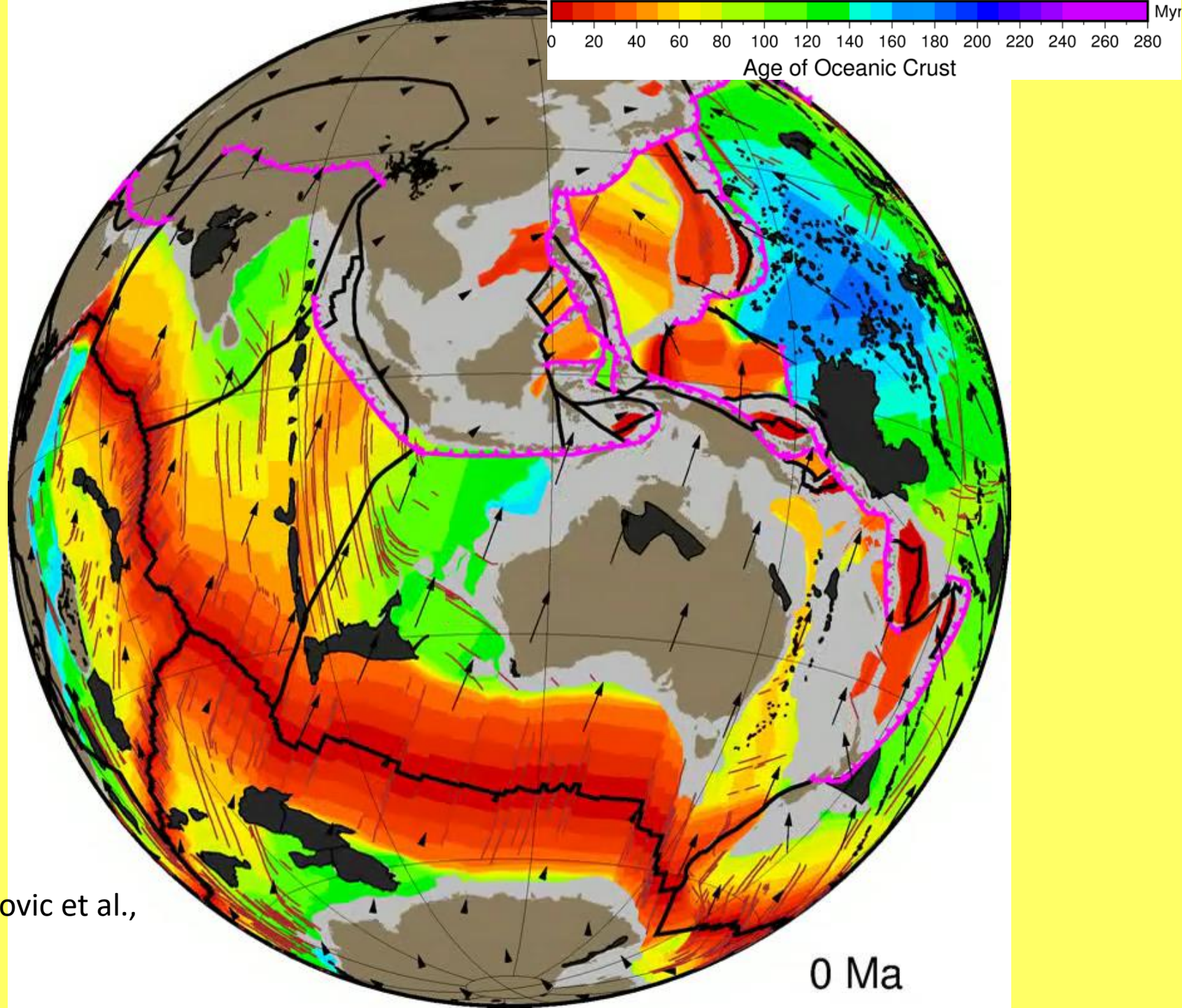
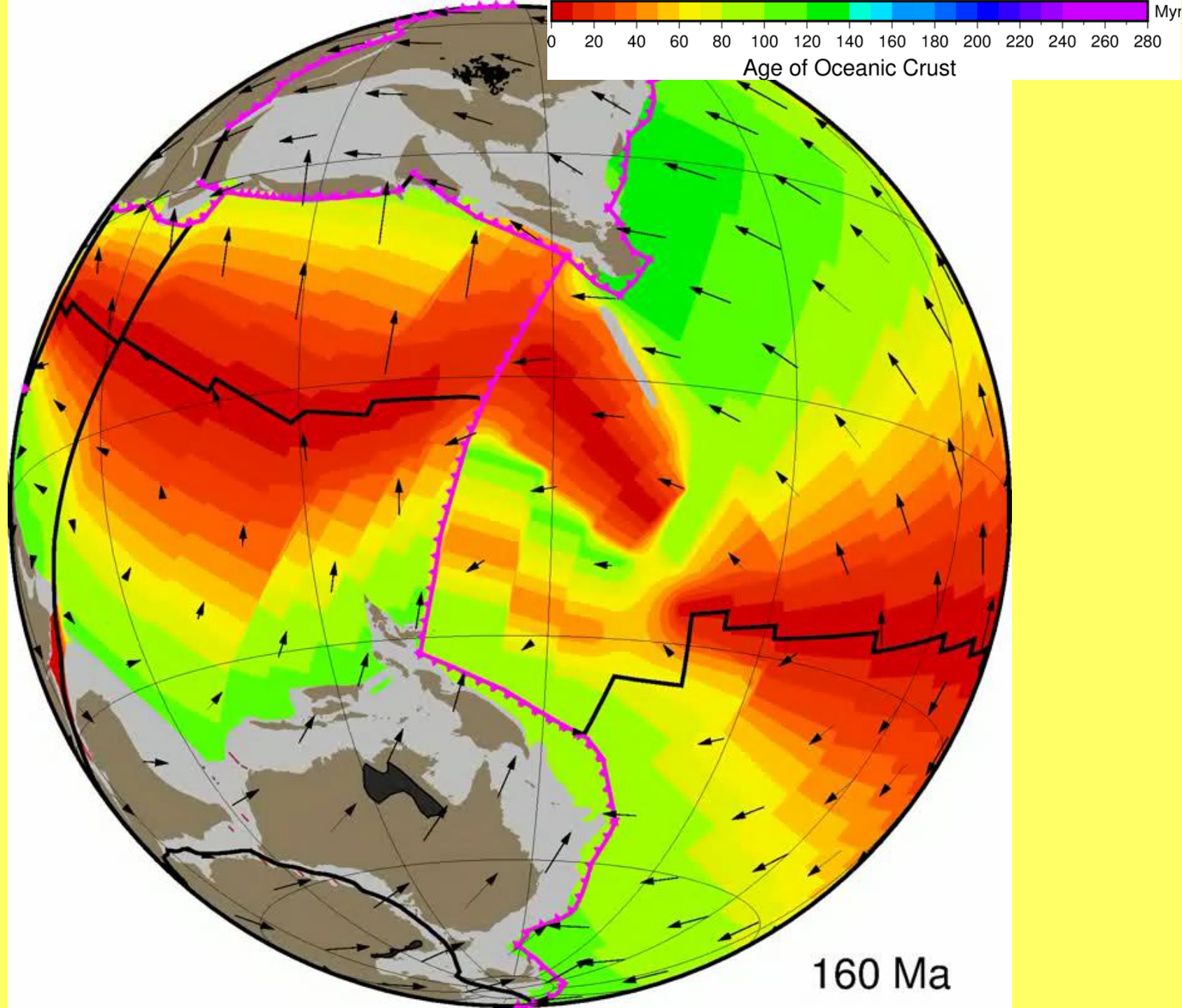
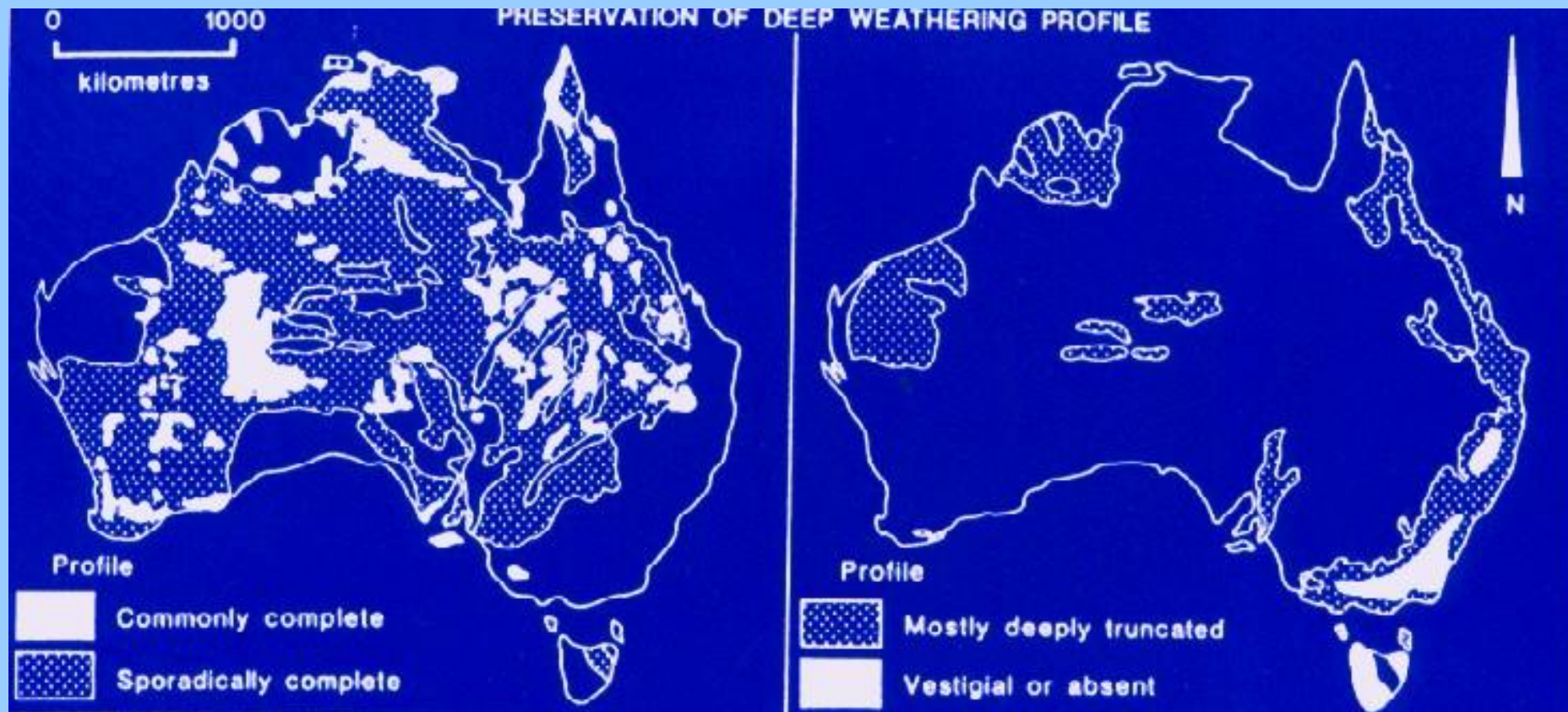


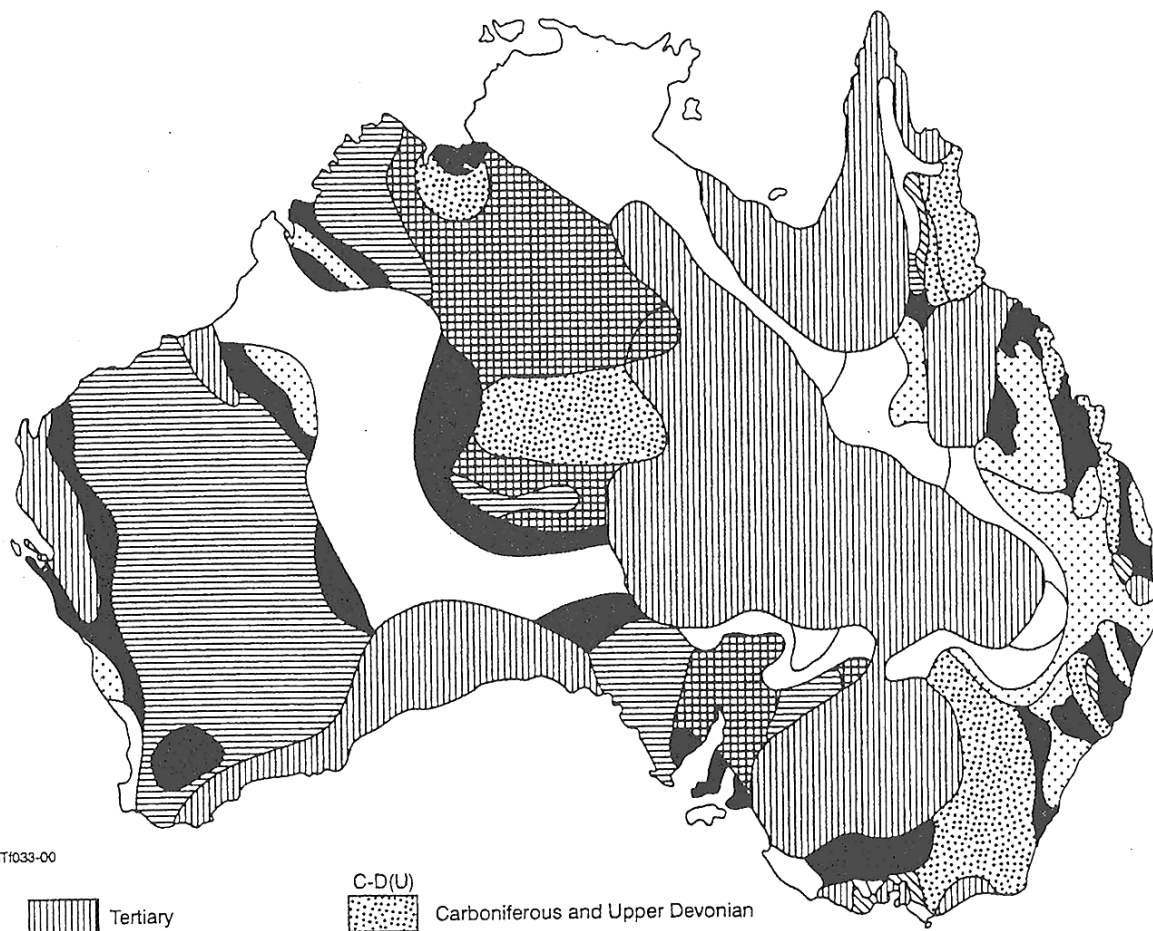
Figure 6. Australia's migration north since the Late Cretaceous (after Anand, 2005, fig. 5)



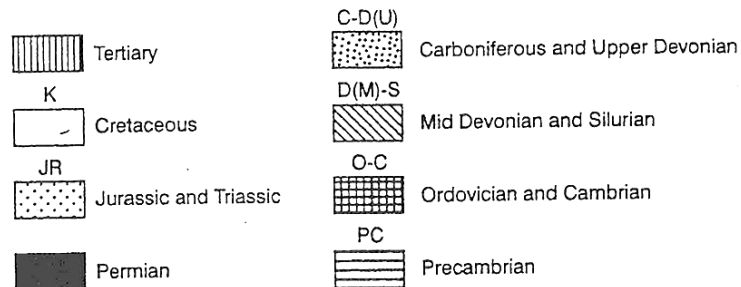
From Zahirovic et al.,
2016








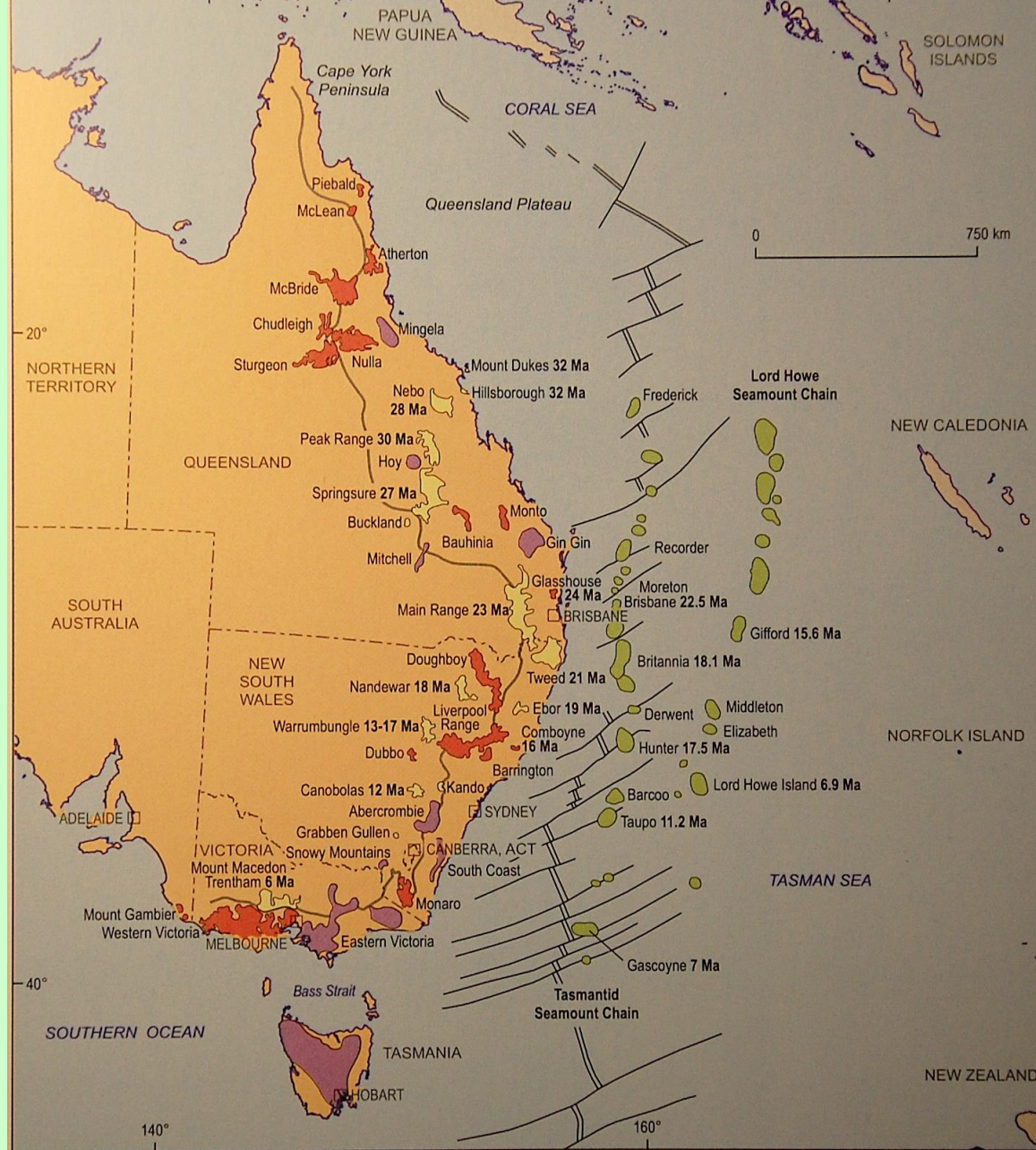
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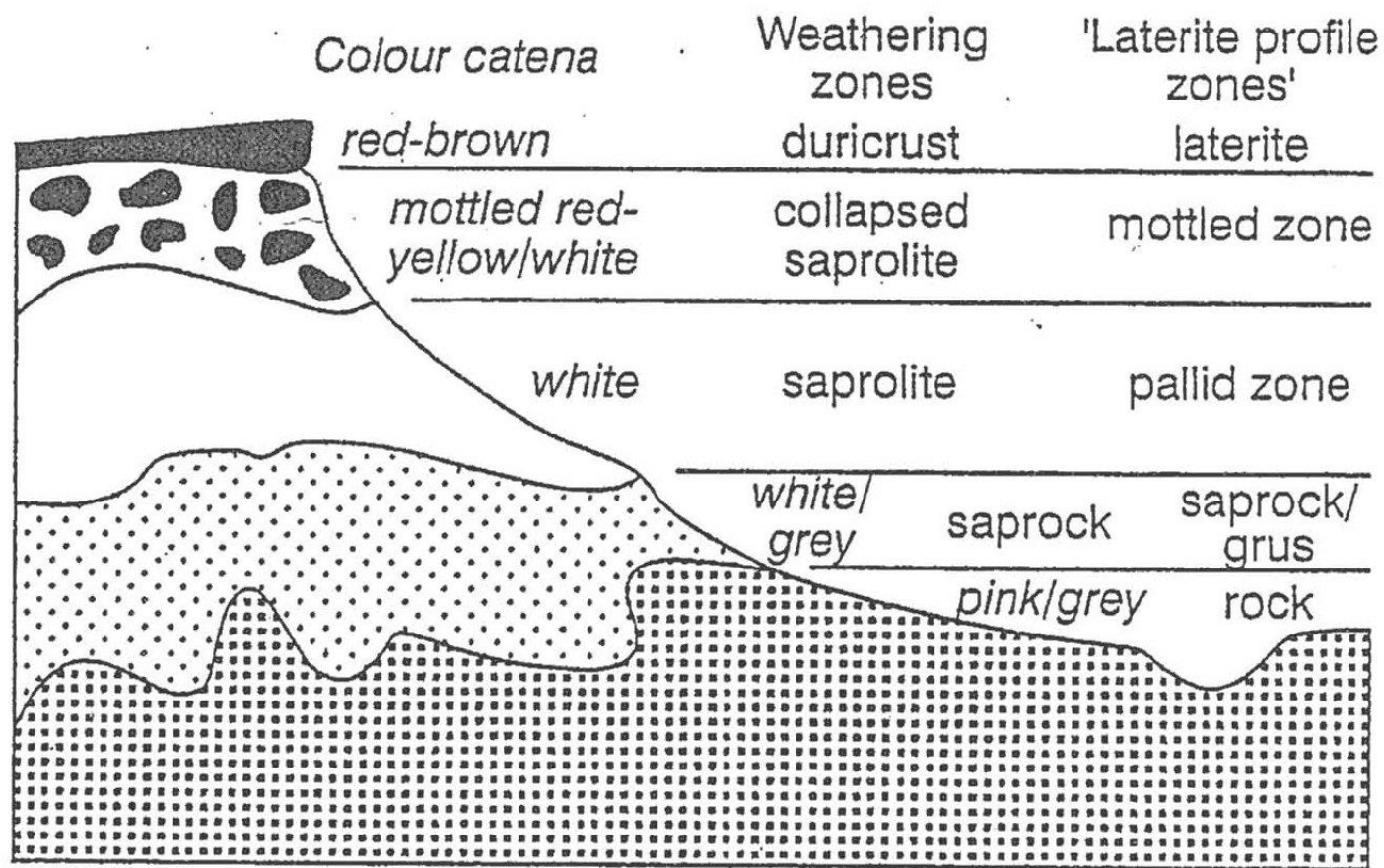


e.g.

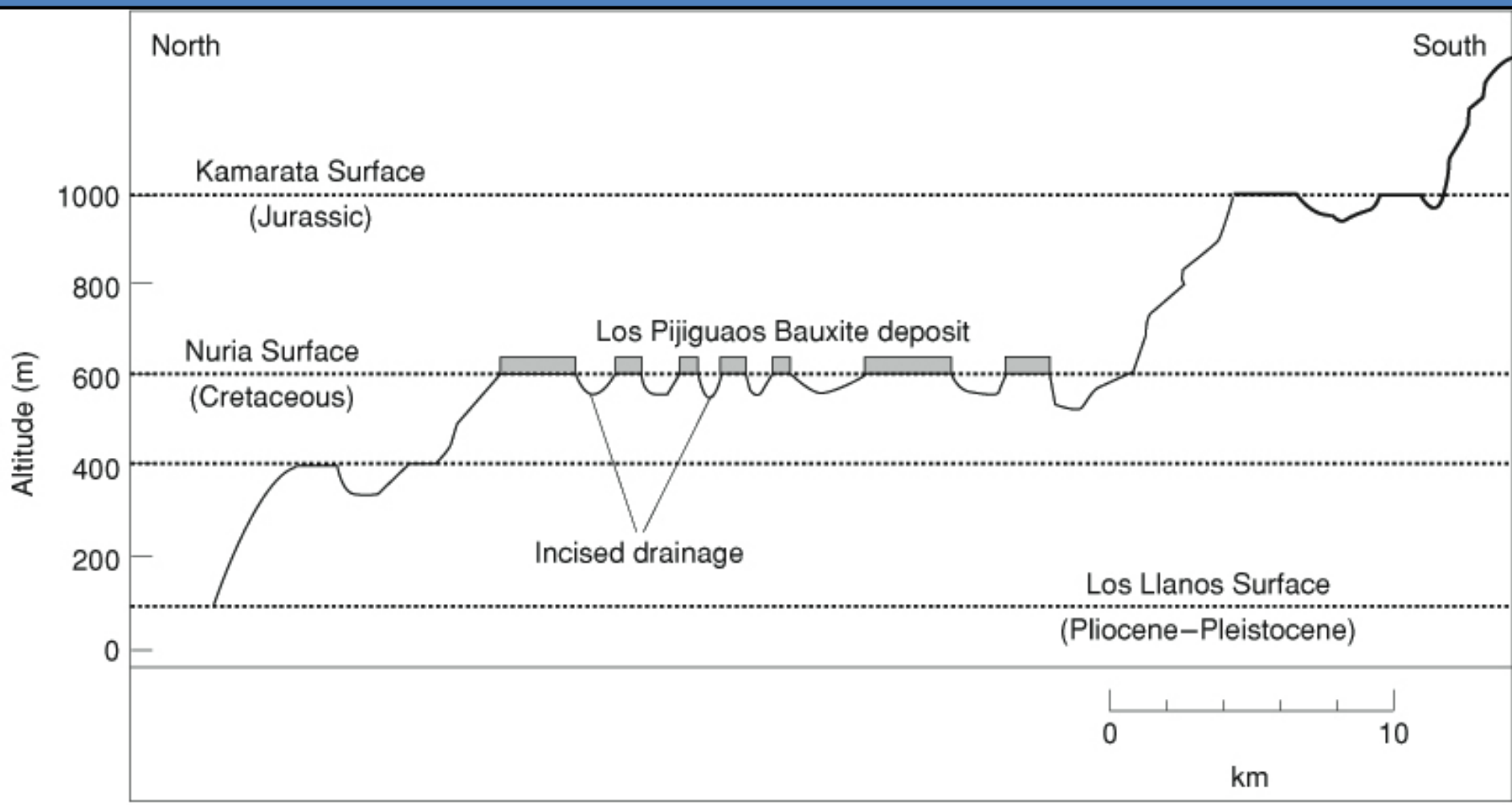
 All of these areas have been continuously exposed to weathering and denudation since the Permian





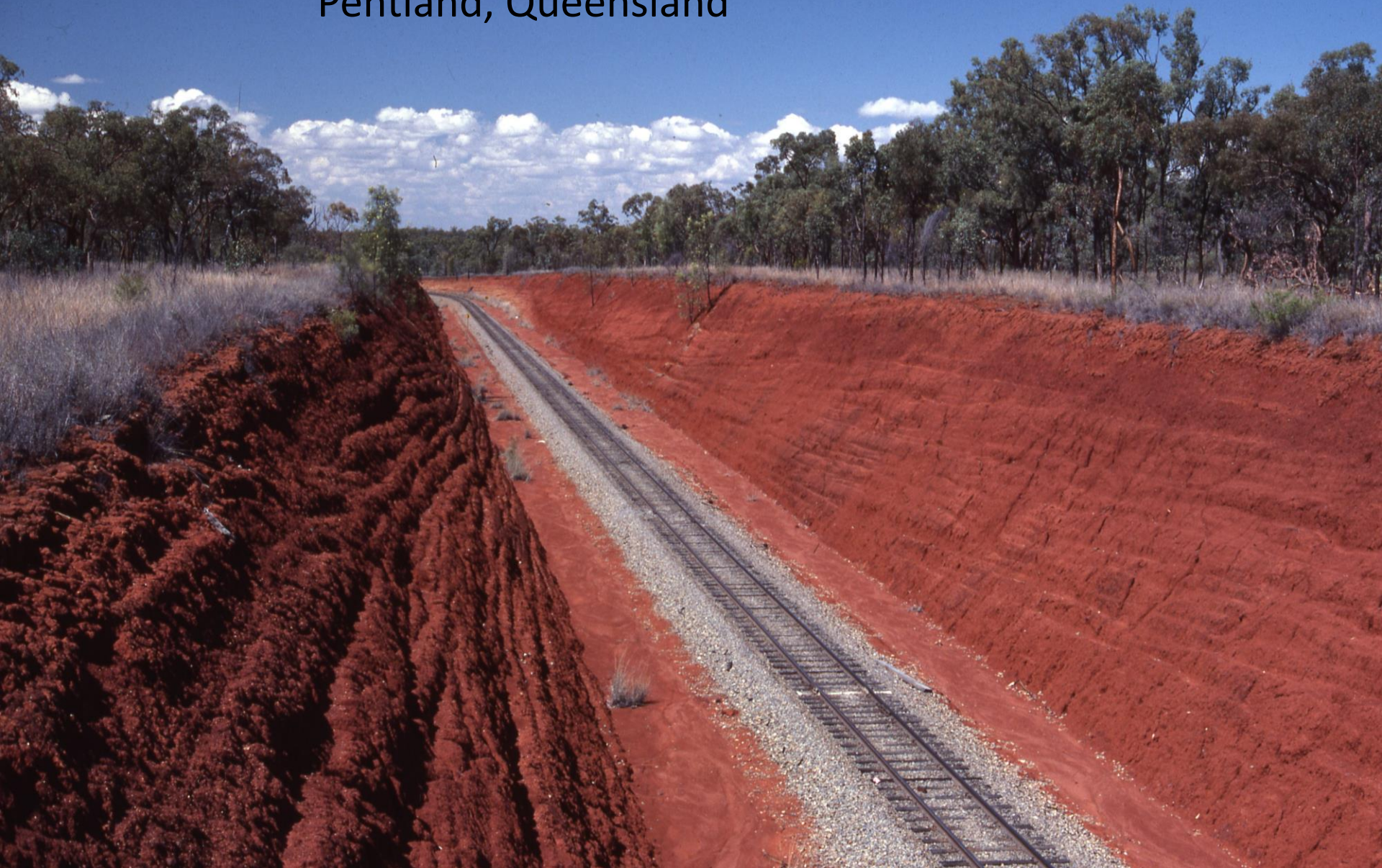


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Pentland, Queensland







Weipa, Far North Queensland (bauxite)



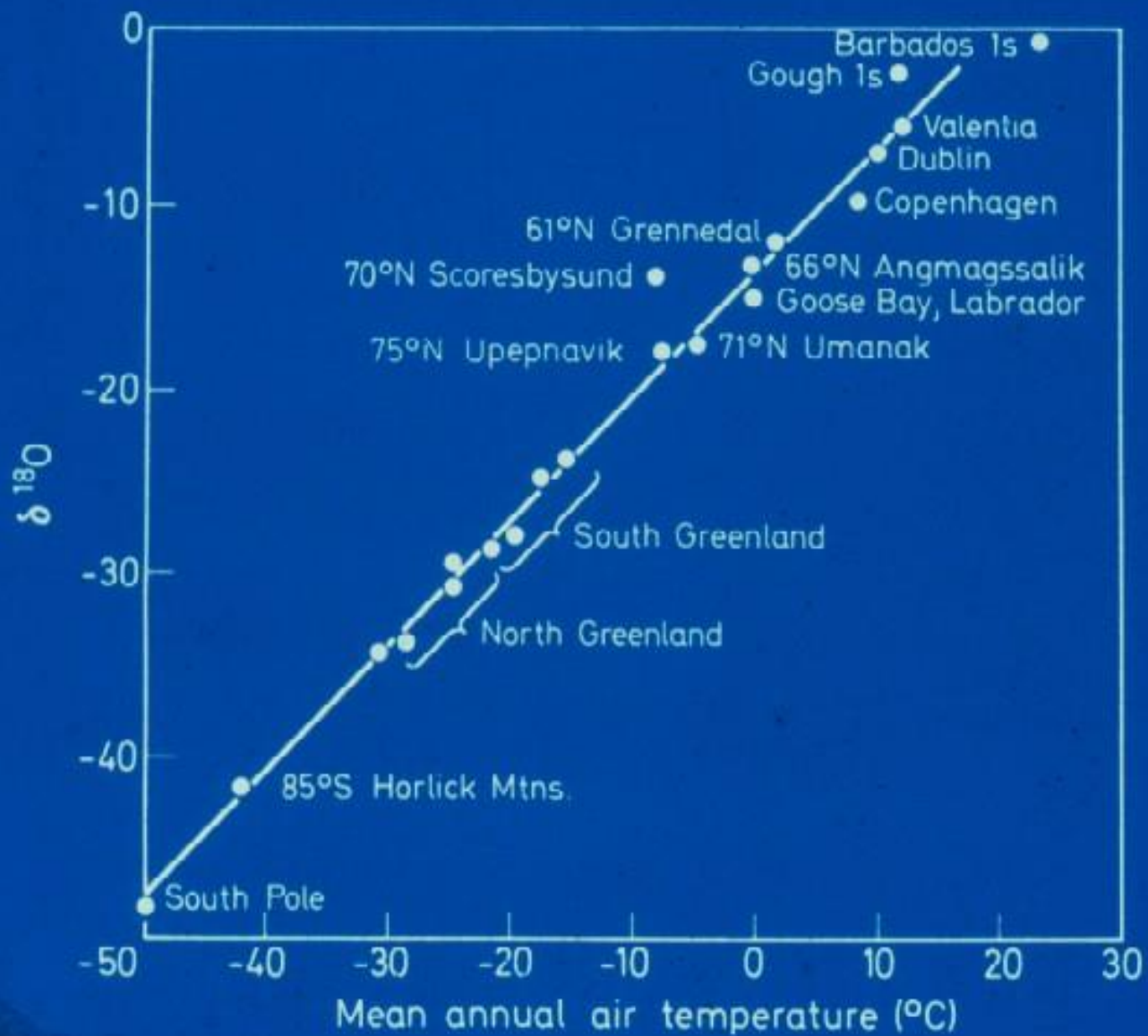


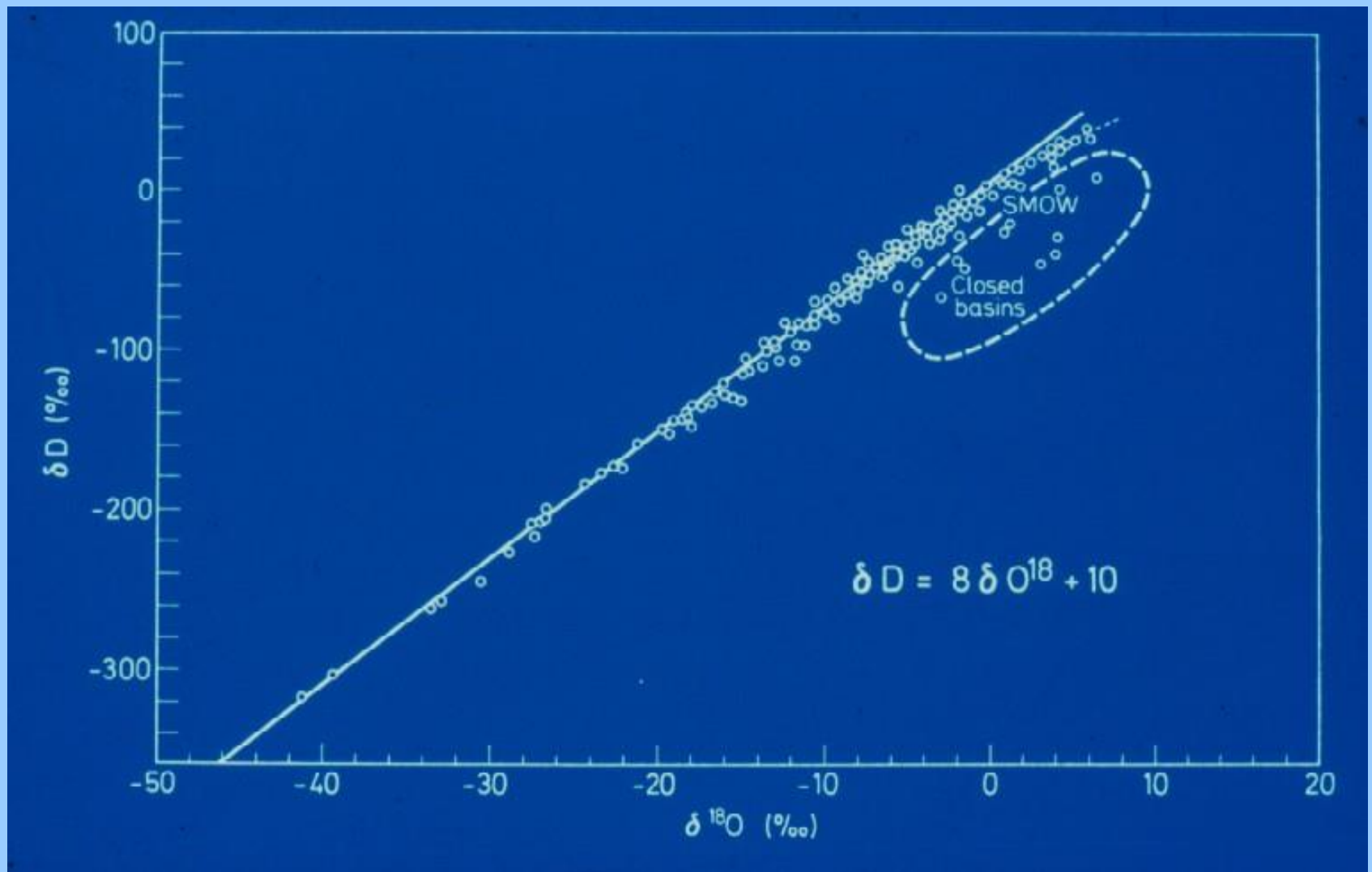


The Breakaways, Coober Pedy, South Australia

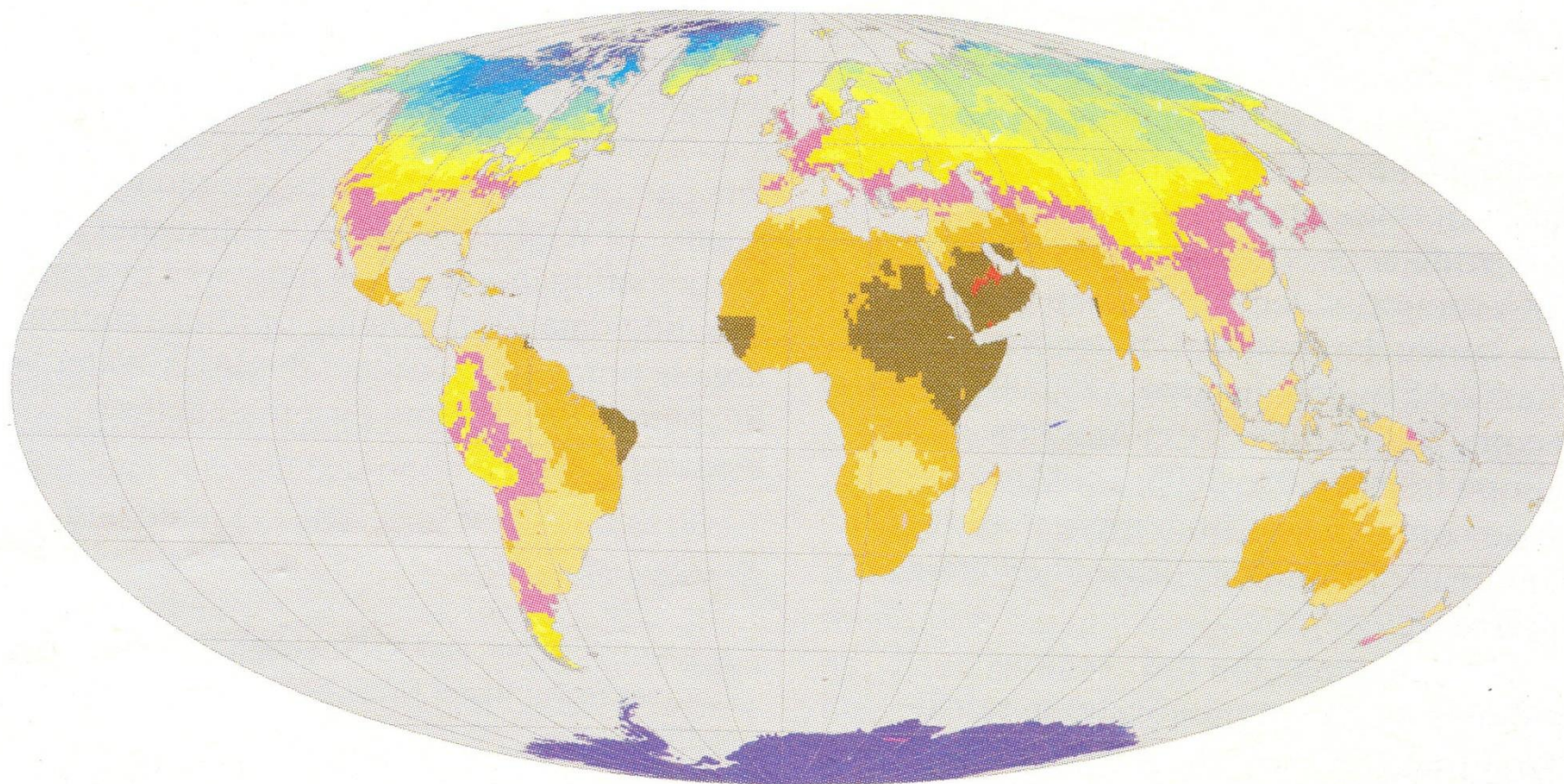


$$\delta^{18}\text{O}_{\text{SMOW}} \text{ (in } \text{‰} \text{)} = \left(\frac{{}^{18}\text{O}/{}^{16}\text{O}_{\text{sample}} - {}^{18}\text{O}/{}^{16}\text{O}_{\text{SMOW}}}{{}^{18}\text{O}/{}^{16}\text{O}_{\text{SMOW}}} \right) 1000$$

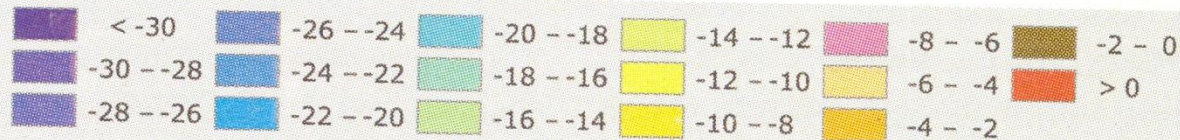


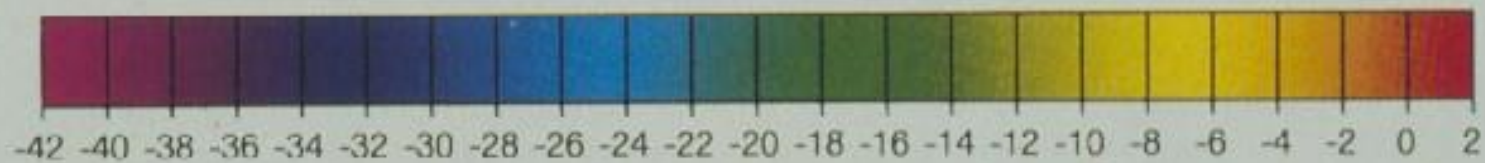
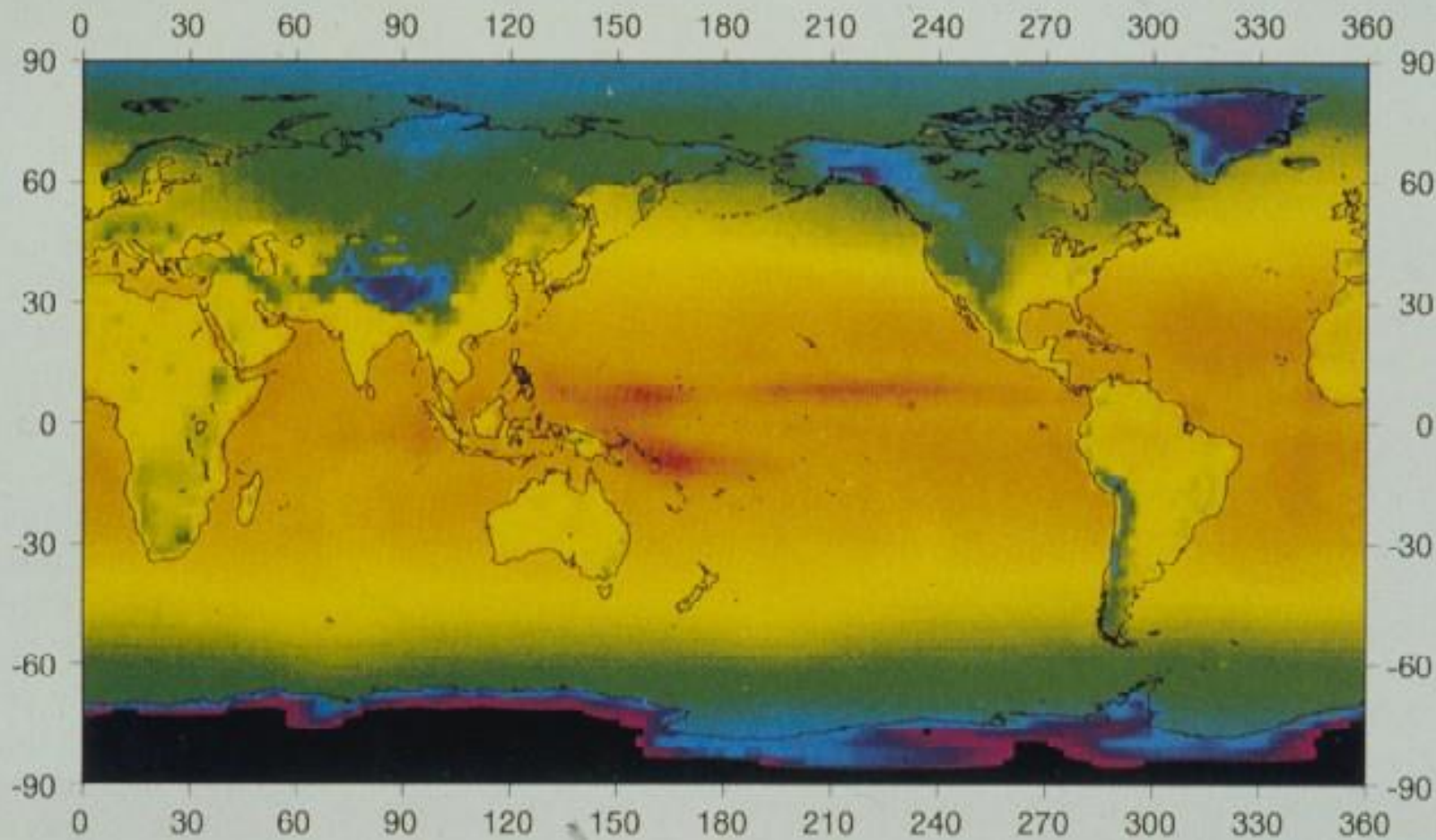


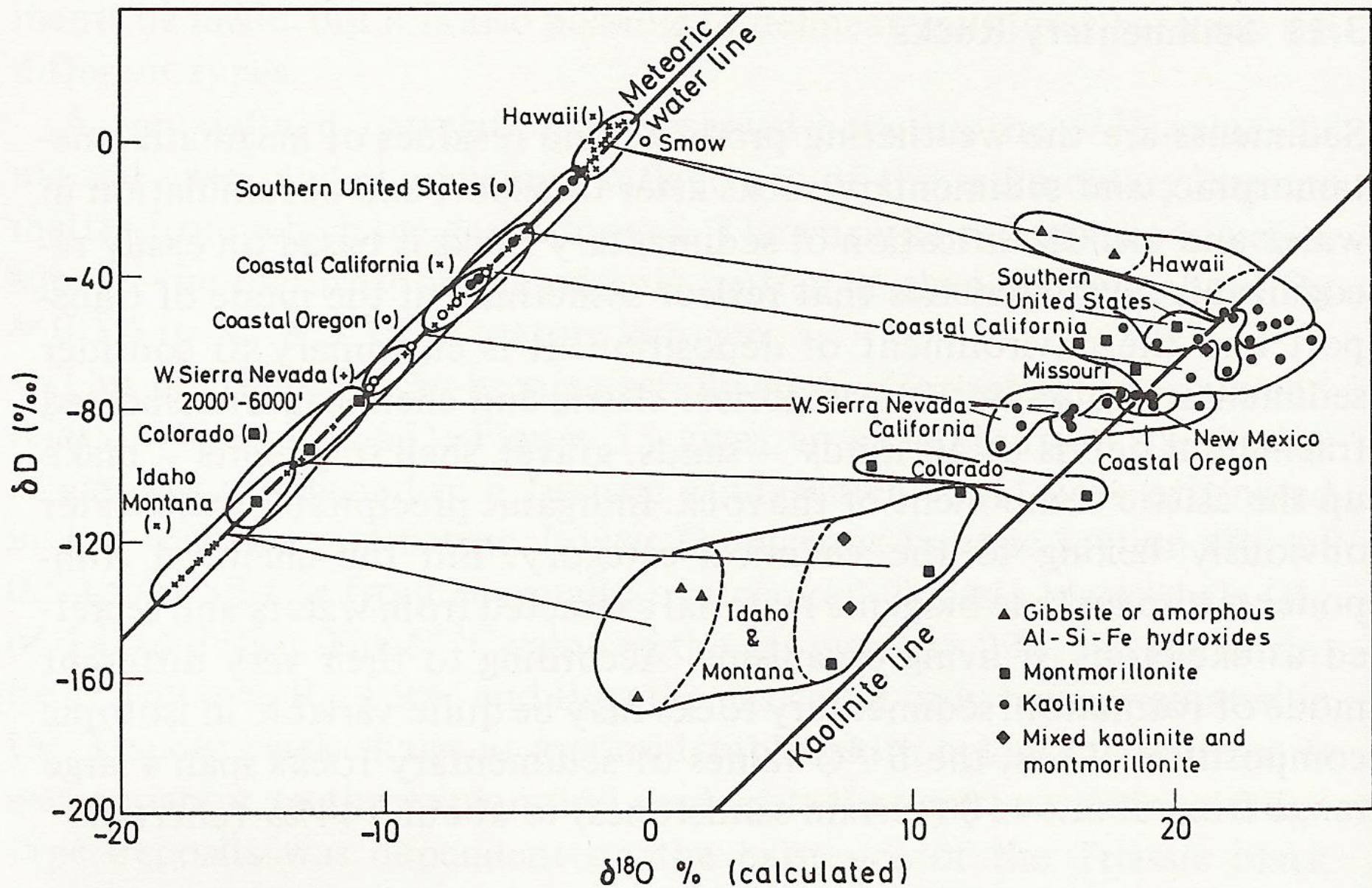
Craig, 1961



$\delta^{18}\text{O}(\text{‰})$

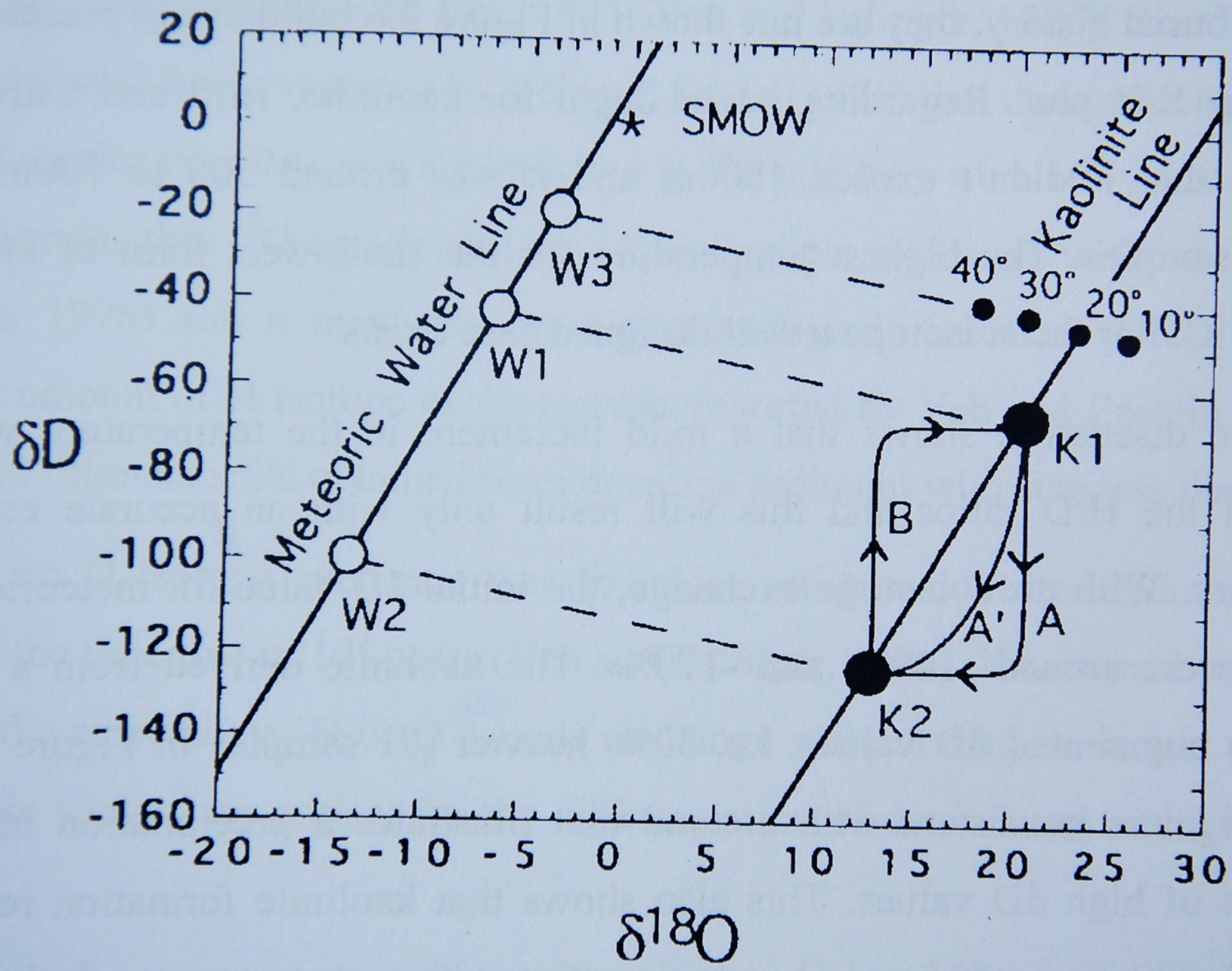






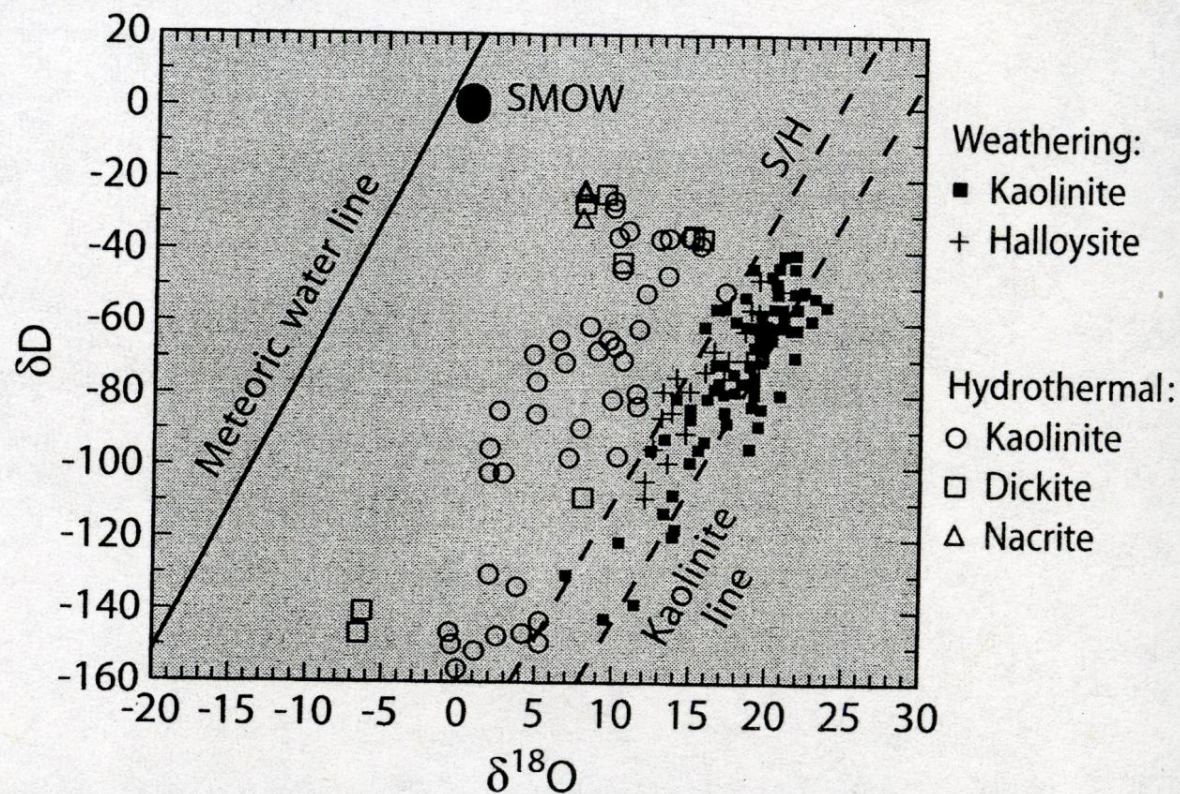
Lawrence and Taylor, 1971

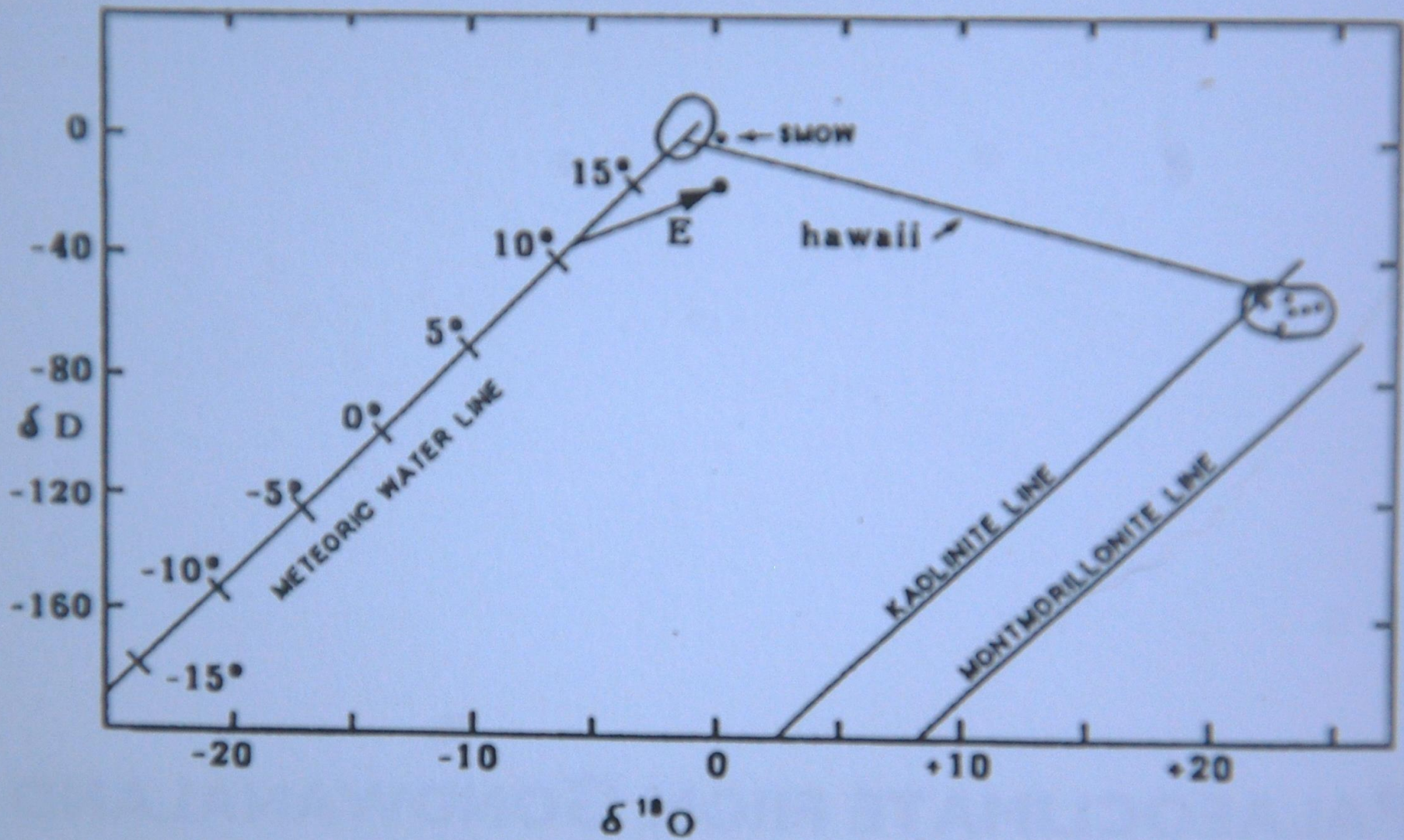
Oxygen fractionation is +27 per mil



(Sheppard and Gilg, 1996)

Fig. 65. Compilation of δD and $\delta^{18}O$ values for kaolin-ites and related minerals from weathering and hydro-thermal environments. The Meteoric Water Line, kaoli-nite weathering line, and supergene/hypogene (S/H) lines are given for reference. (After Sheppard and Gilg 1995)



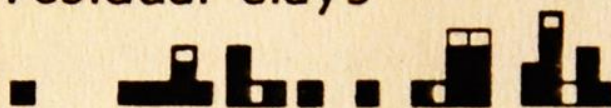


(Chivas and Bird, 1995)

Post mid-Tertiary residual clays



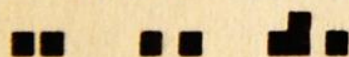
Pre mid-Tertiary residual clays



Pre late-Mesozoic residual clays



Permian residual clays



+7

+9

+11

+13

+15

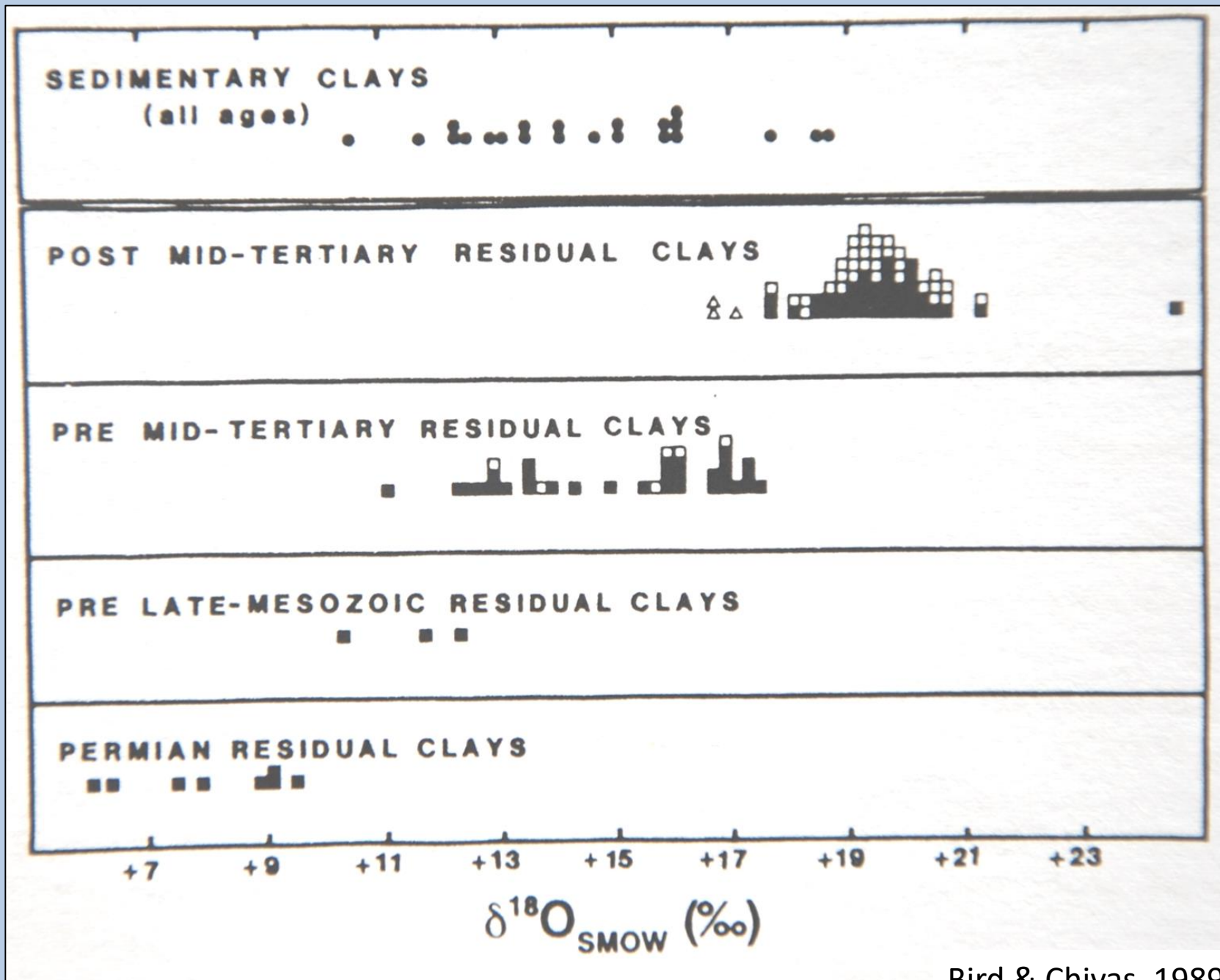
+17

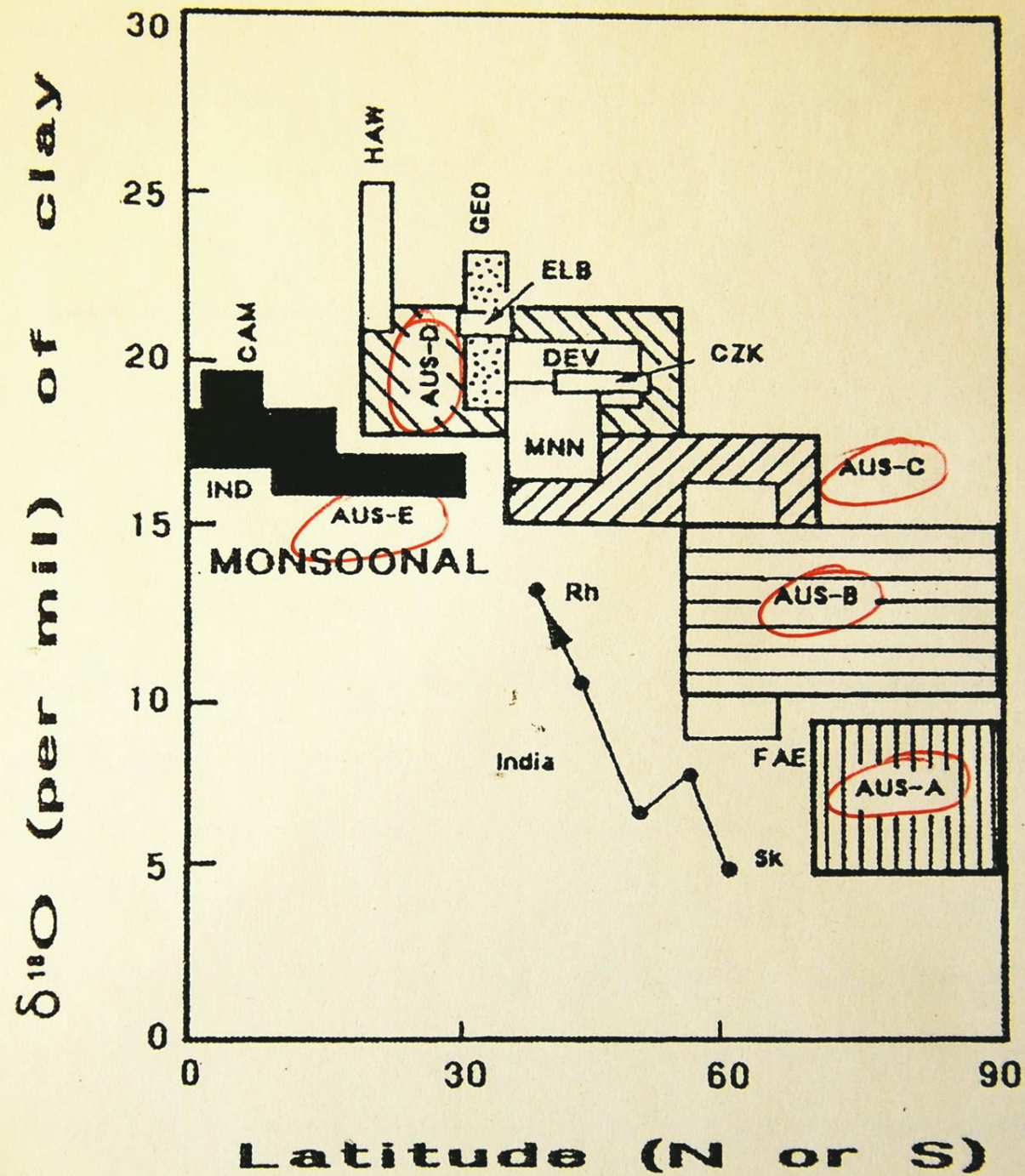
+19

+21

+23

$\delta^{18}\text{O}_{\text{SMOW}}$ (‰)



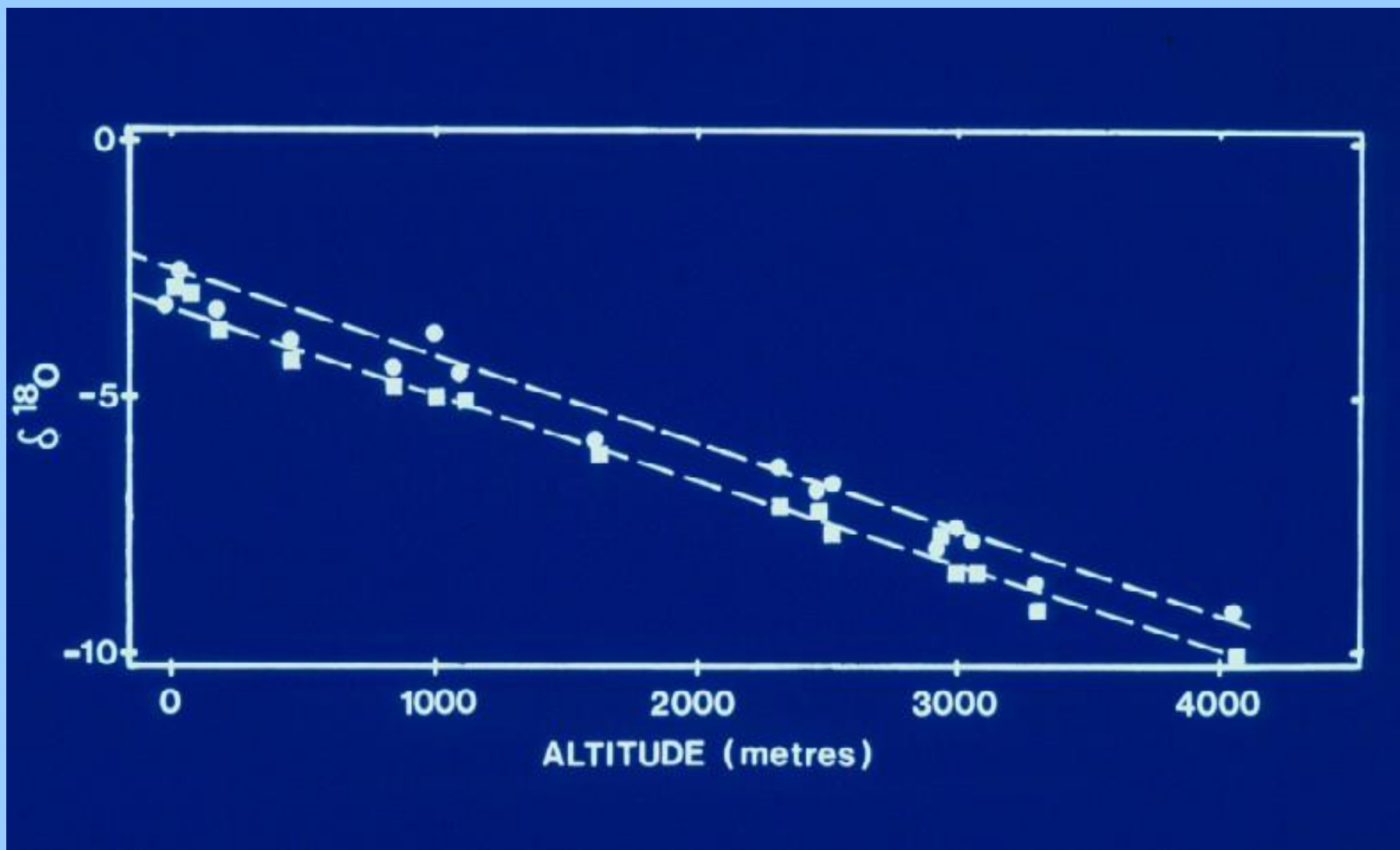


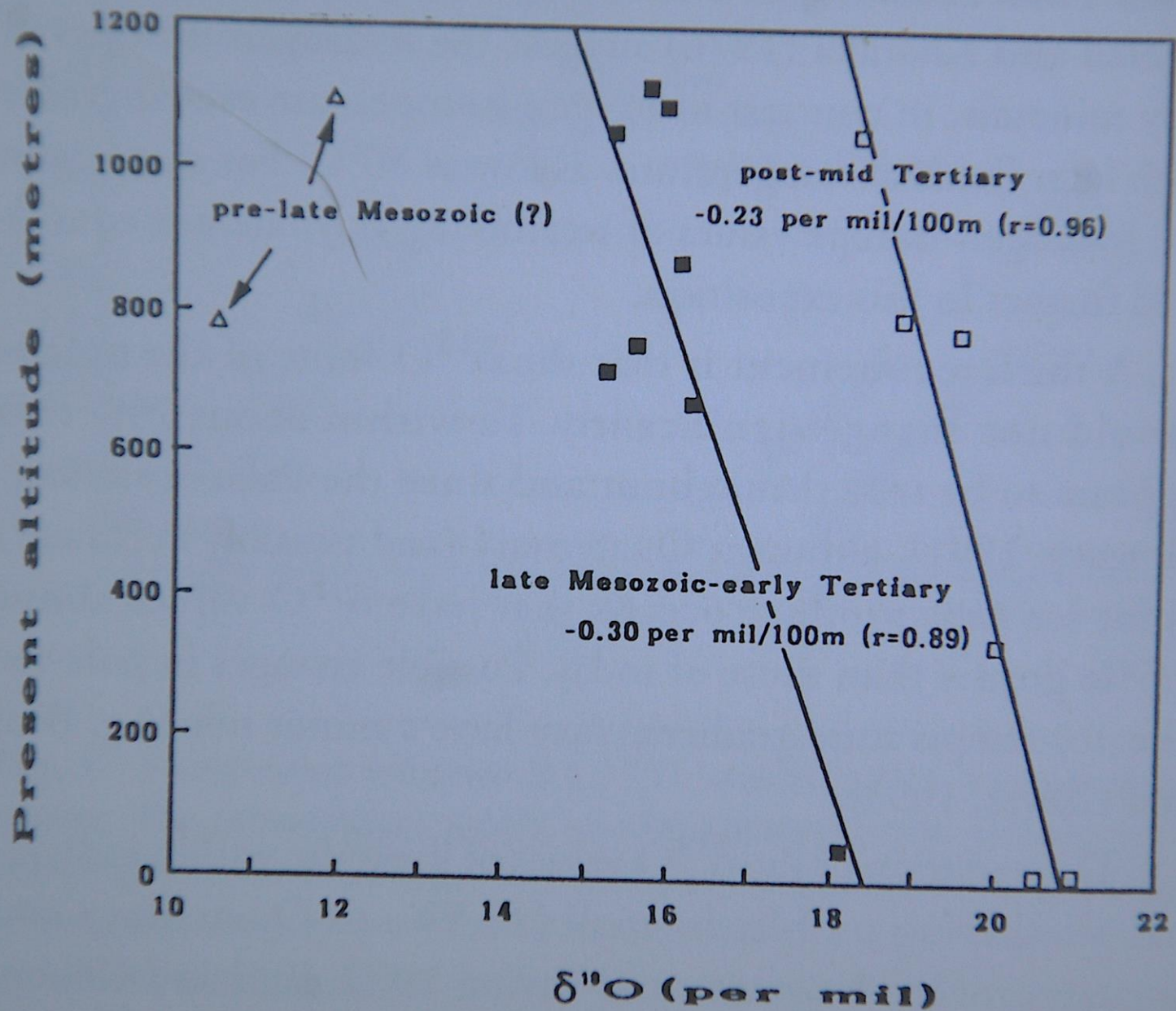
Principal Conclusions of $\delta^{18}\text{O}$ Studies

Excluding monsoonal regions to the north, it is possible to distinguish profiles formed in –

- Late Mesozoic – Early Cenozoic = $\delta^{18}\text{O}$ +15 to +17.5‰
- Post-mid-Cenozoic weathering profiles = $\delta^{18}\text{O}$ >+17.5 ‰

Also identified remnants of earlier phases of deep weathering (pre-late Mesozoic: Early Cretaceous or Jurassic?) with $\delta^{18}\text{O}$ +10 to +15‰ indicating weathering in cool to cold and presumably humid climate





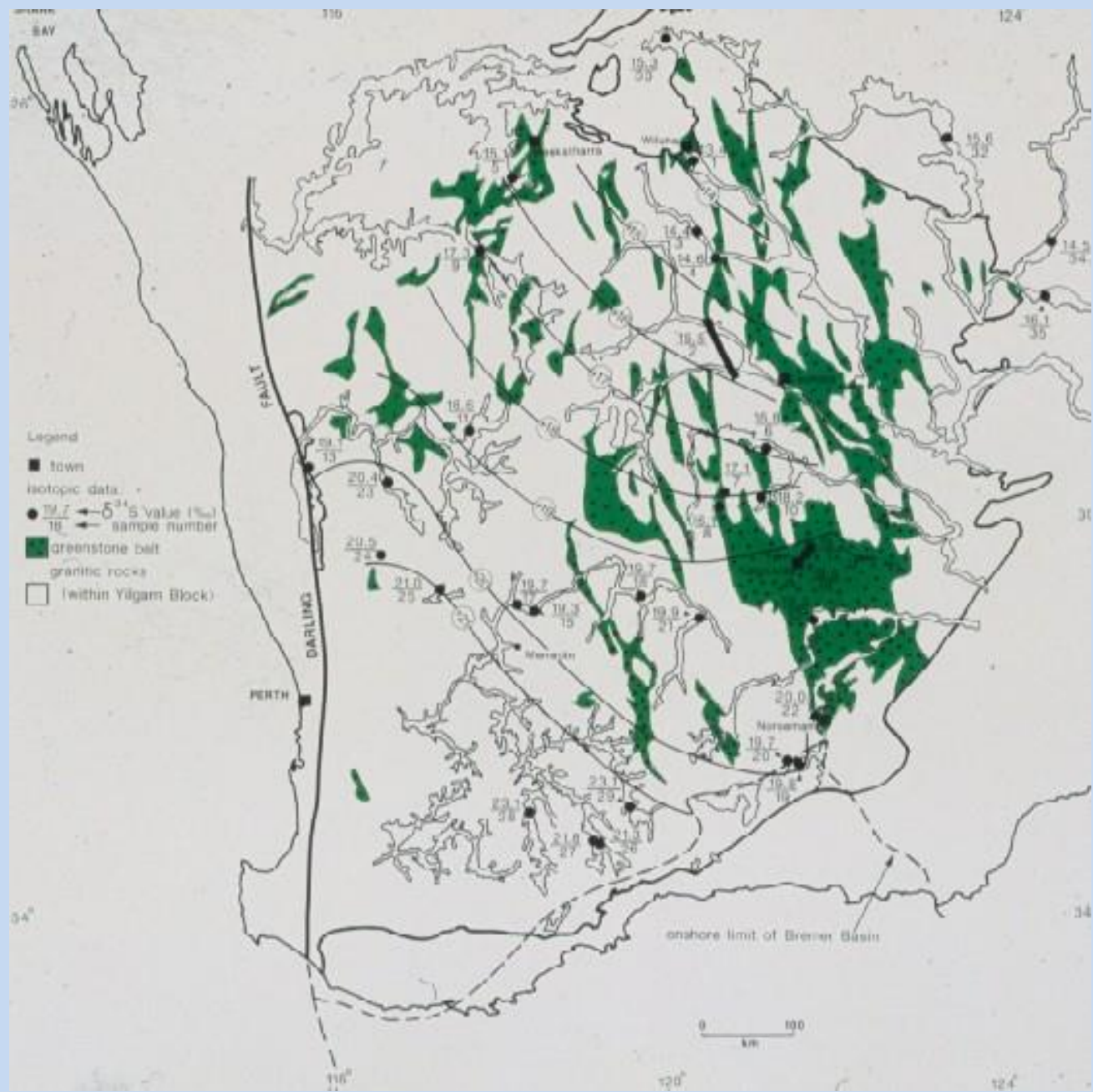
(Chivas and Bird, 1995)

Stable Isotope Studies and the Timing of Weathering Events

Bird & Chivas (1993) *Australian Journal of Earth Sciences*, **40**, 345-358

$\delta^{18}\text{O}$ ($^{18}\text{O}/^{16}\text{O}$) measured in authigenic clay minerals as a parameter of weathering (clay formation) age (isotopic fractionation)

$\delta^{18}\text{O}$ ratio reflects the changing isotopic composition of the meteoric waters in response to changes brought about by the northward movement of the continent since middle Cretaceous time (c. 95 Ma)



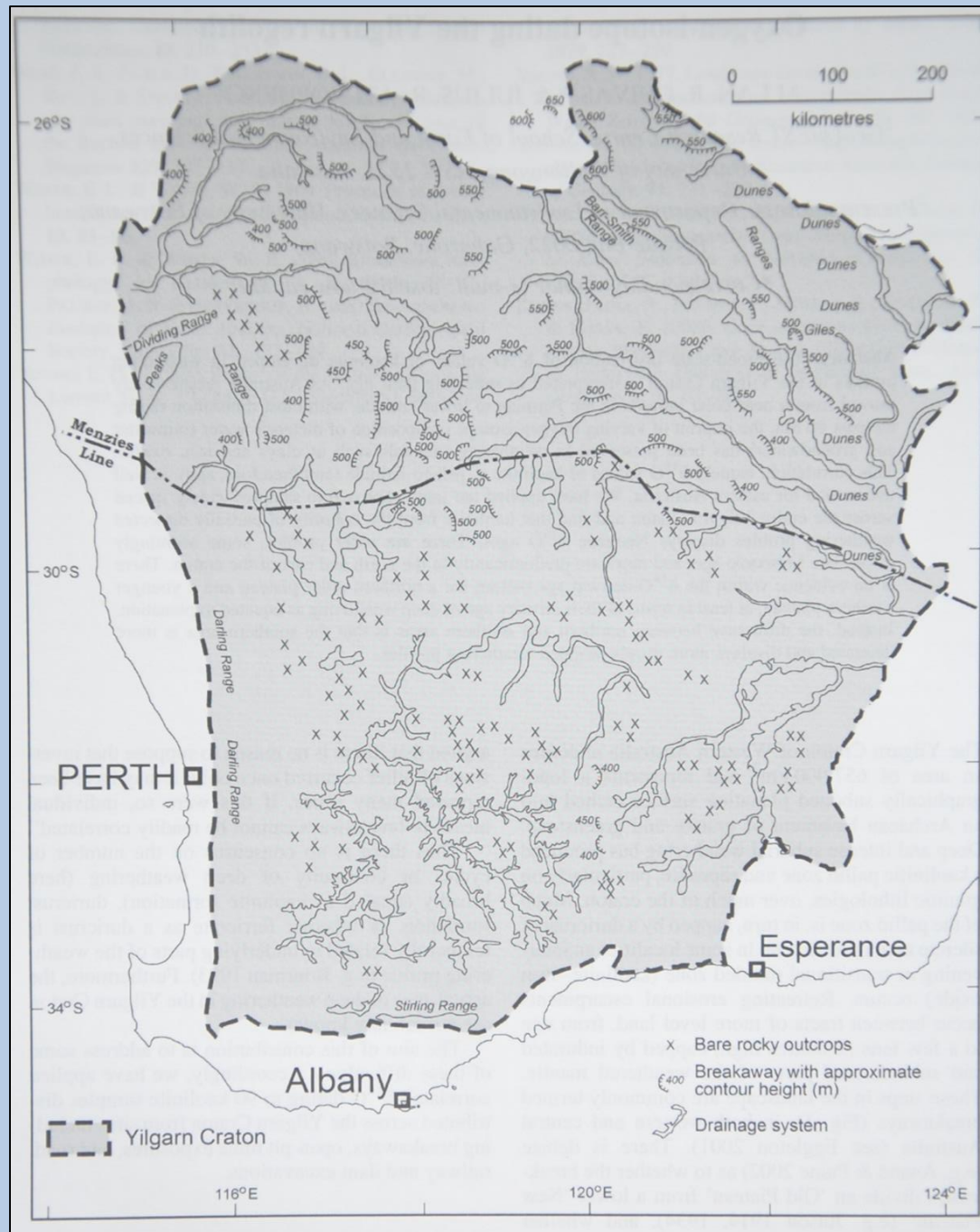


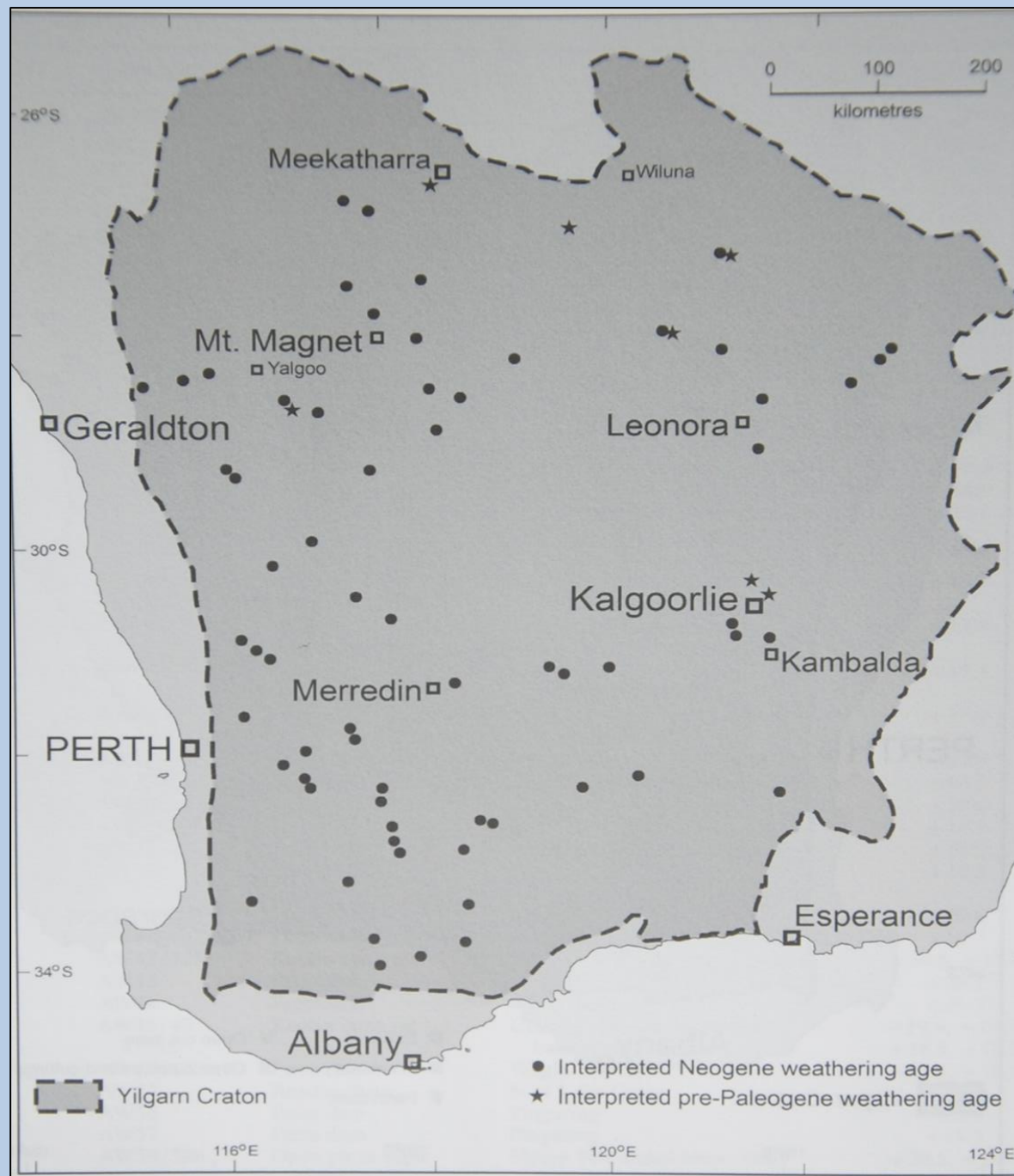
Hyden-Norseman Road





Merredin, Western Austral. Granite substrate, Yilgarn Craton (Archaean)





The range of $\delta^{18}\text{O}$ values for kaolinite from weathering profiles from the Yilgarn craton:

Most are in the range +22 to +17 and are indicative of a Neogene age.

Some are in the range +15 to +17 and are plausibly of Late Cretaceous to Palaeogene age
(Kalgoorlie, Lawlers, Mt Percy, Bluff Point)

A few have even lower values:

Scuddles, south of Yalgoo (+14.6 per mil)

Bronzewing, Discovery pit (+14.5 per mil)

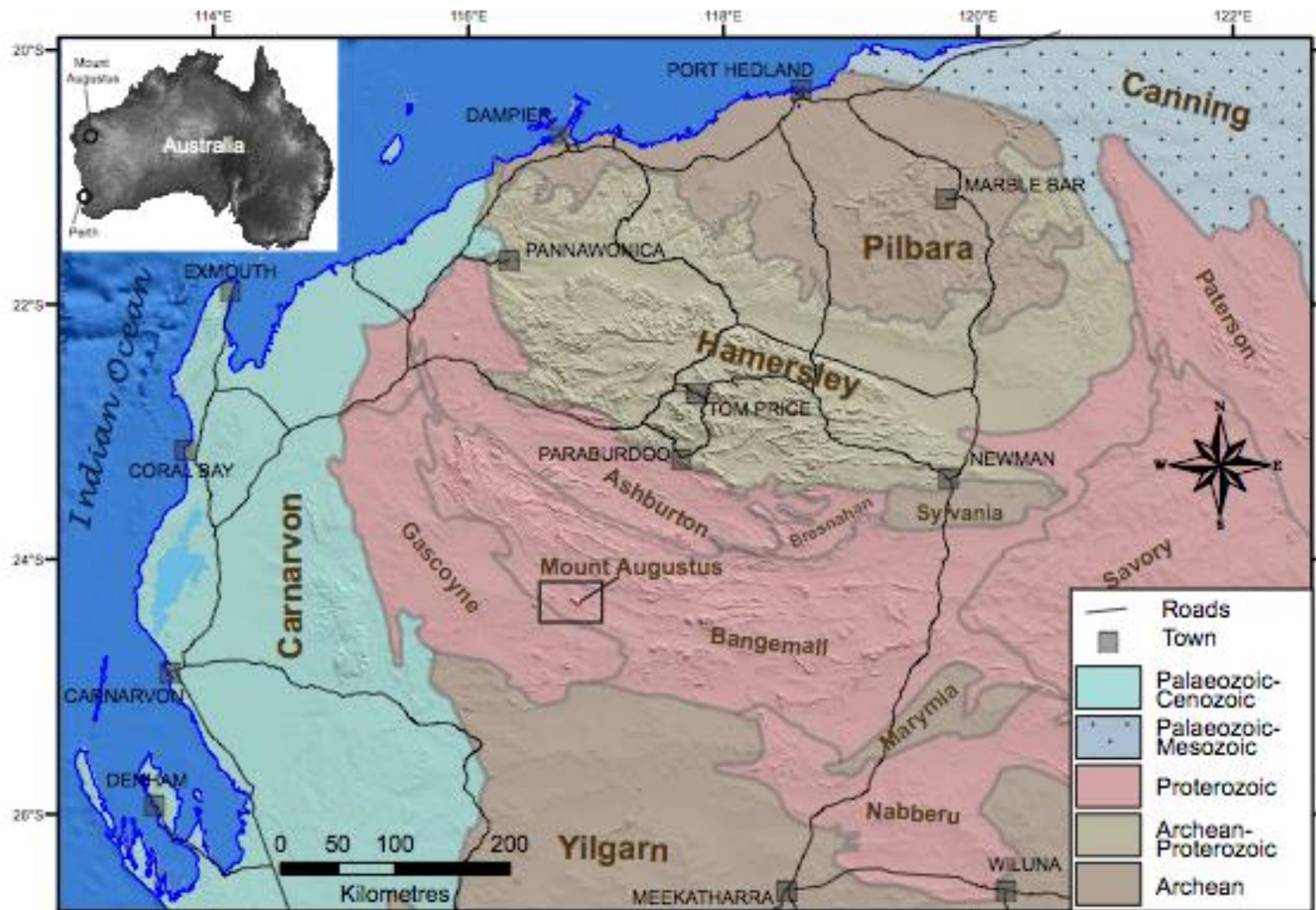
Meekatharra (+13.2 per mil)

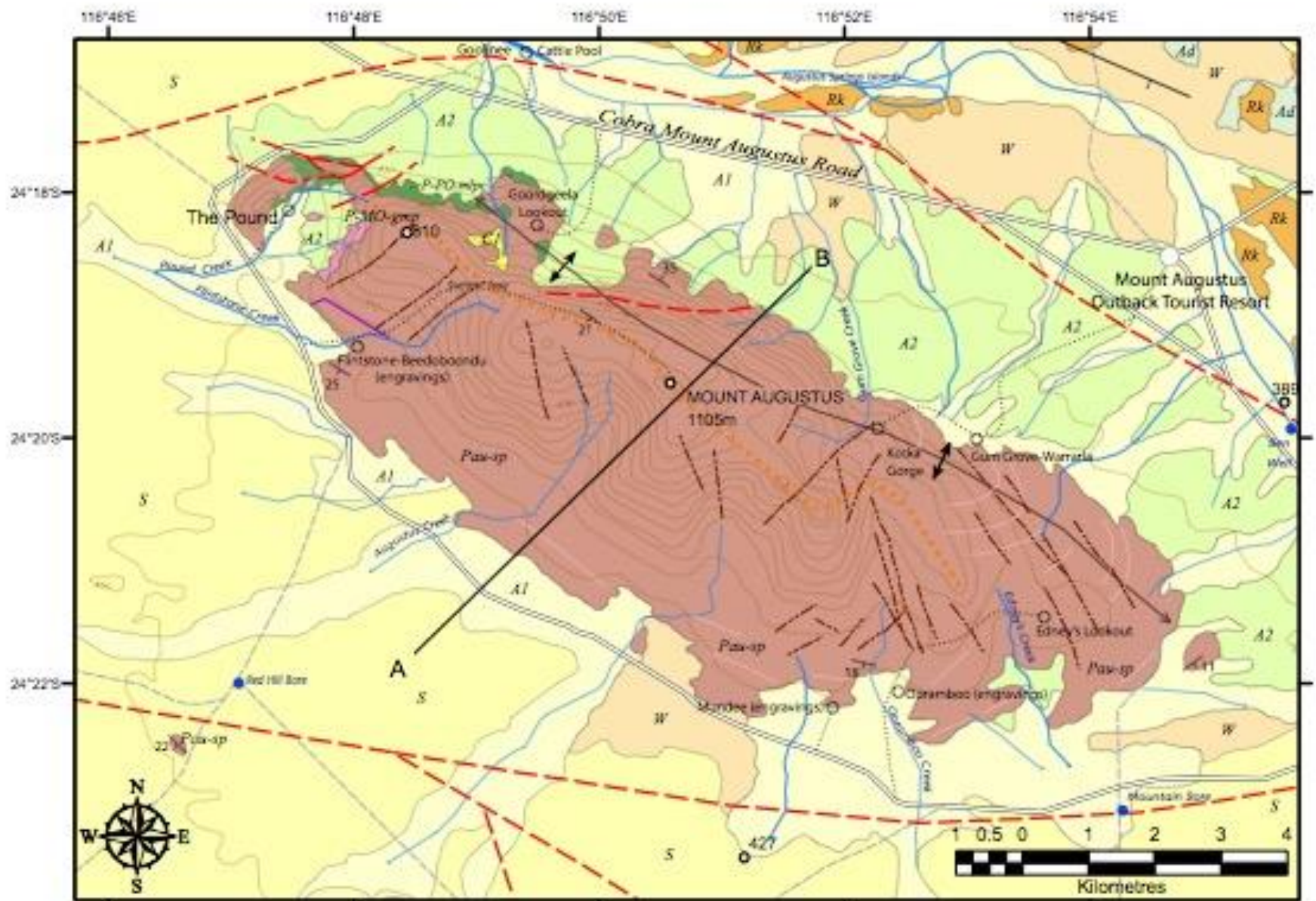
and are also of 'old' age, either as above or potentially older (Jurassic?)

No known Permian kaolinites were measured (none was sampled)

Burringurrah, or Mount Augustus, Western Australia
(15 km)

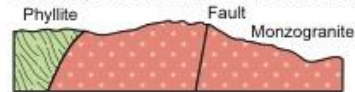




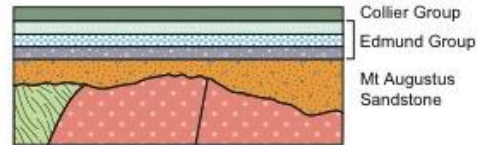




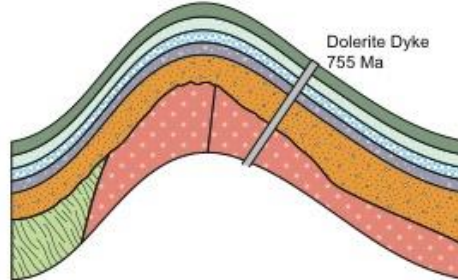
A) Monzogranite of Moorarie Super-Suite (1830-1780 Ma) and Pooranoo Metamorphics (1680 Ma) affected by Mangaroo Orogeny (1680-1620 Ma). Erosion of monzogranite



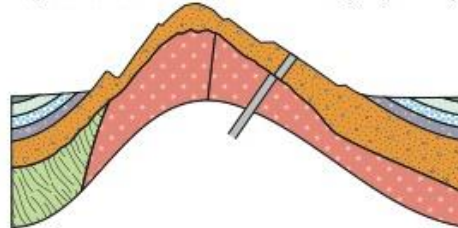
B) Deposition of Mount Augustus Sandstone (~1620 Ma), Edmund Group (~1465 Ma) and Collier Group ~1070 Ma.



C) Folding - Edmundian Orogeny 1030-950 Ma and intrusions of Mundine Well Dolerite Suite dolerite dyke (755 Ma).



D) Erosion of Edmund Group rocks and some of Mount Augustus sandstone. Disruption of dolerite dyke (~570 Ma)



E) Further erosion of Mount Augustus Sandstone exposing monzogranite in 'The Pound'. Monzogranite subsequently weathered.





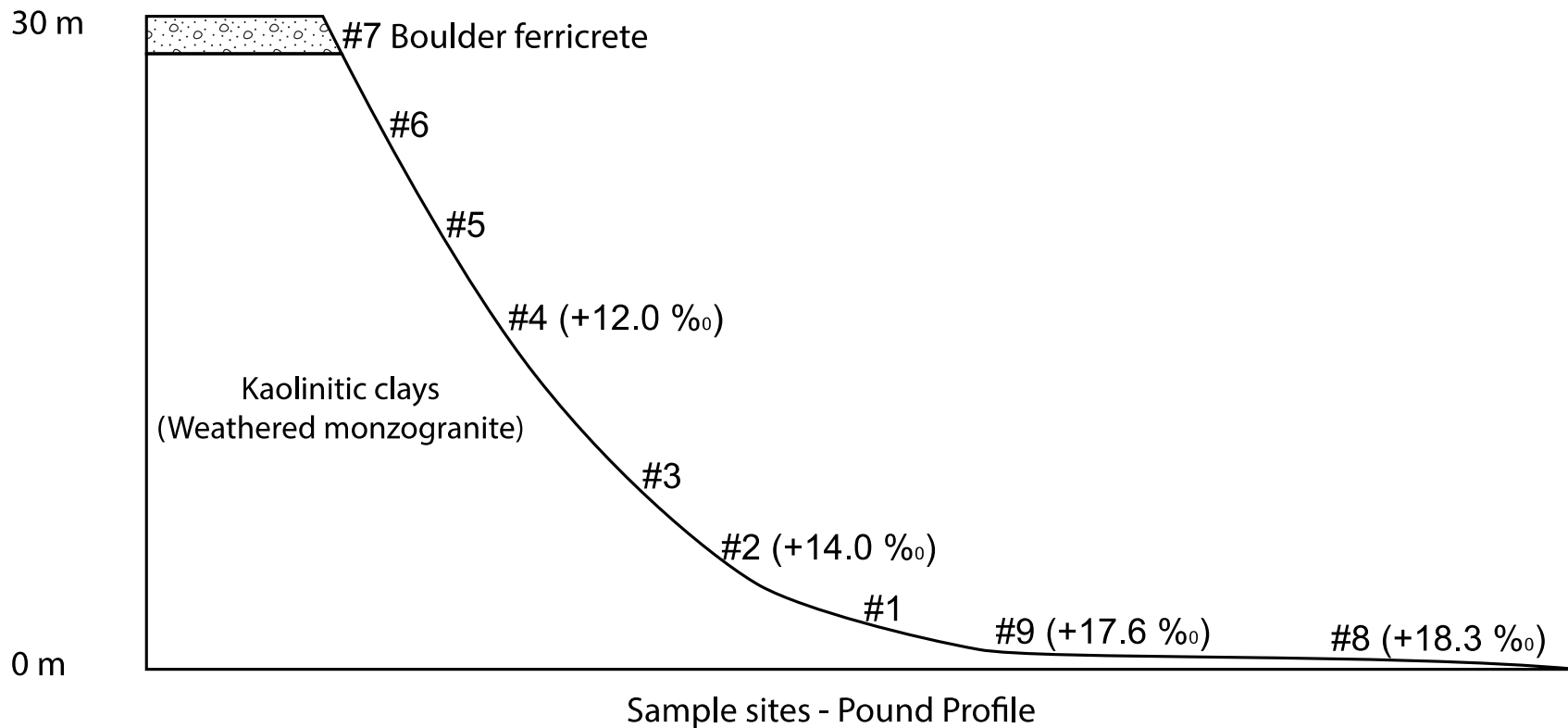












The flank of the Burrungurrah inselberg or monolith has been exposed to sub-aerial weathering for a long period. The progressively deepening weathering profile has the following approximate ages:

- +12.0 and +14.0 per mil = Mesozoic and/or Paleogene
- +17.6 per mil = Paleogene and/or Neogene
- +18.3 per mil = Neogene

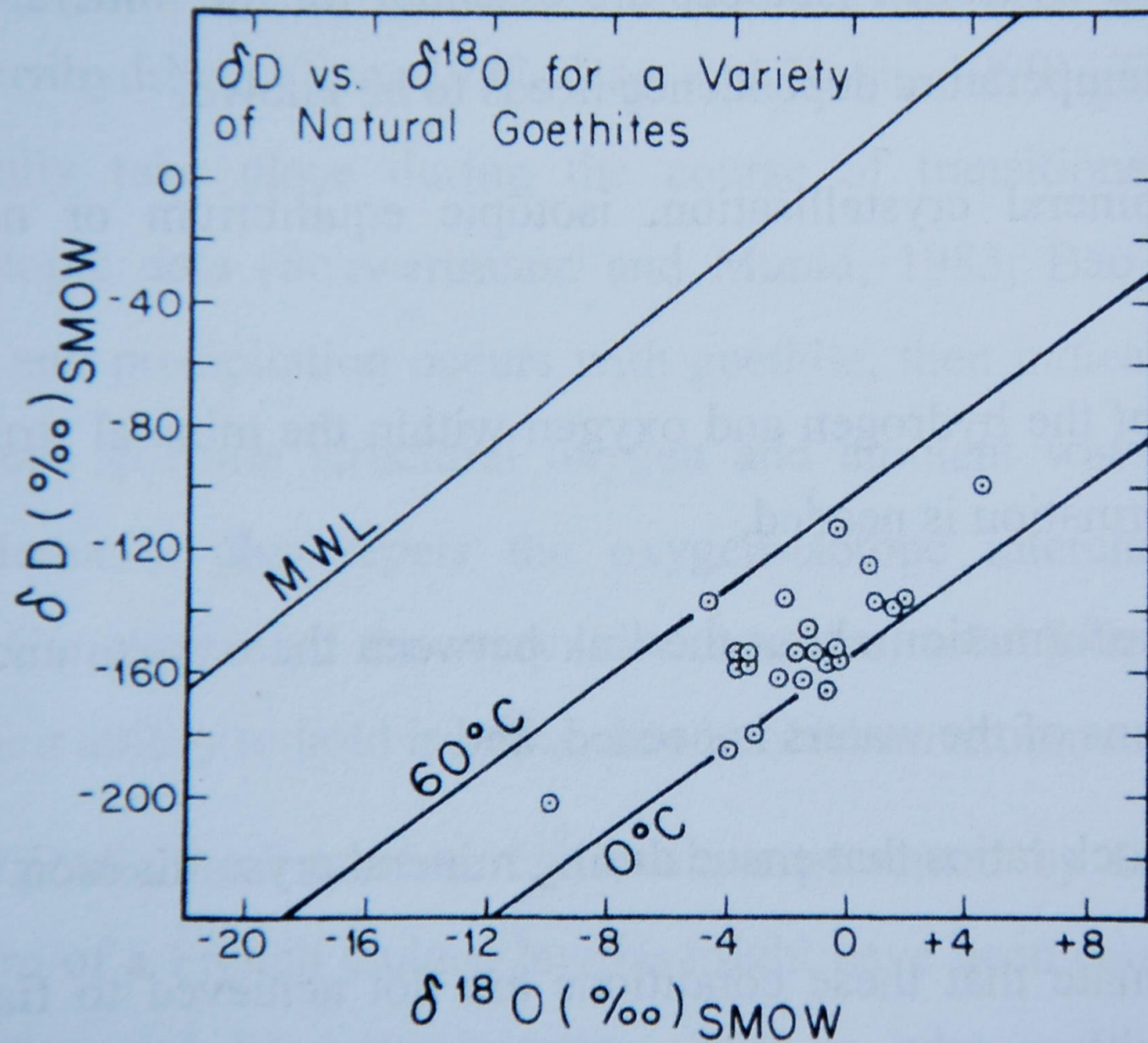
A feature in the landscape for approximately 150 to 100 Ma ?

Deep and intense weathering is not necessarily restricted to 'tropical' latitudes.

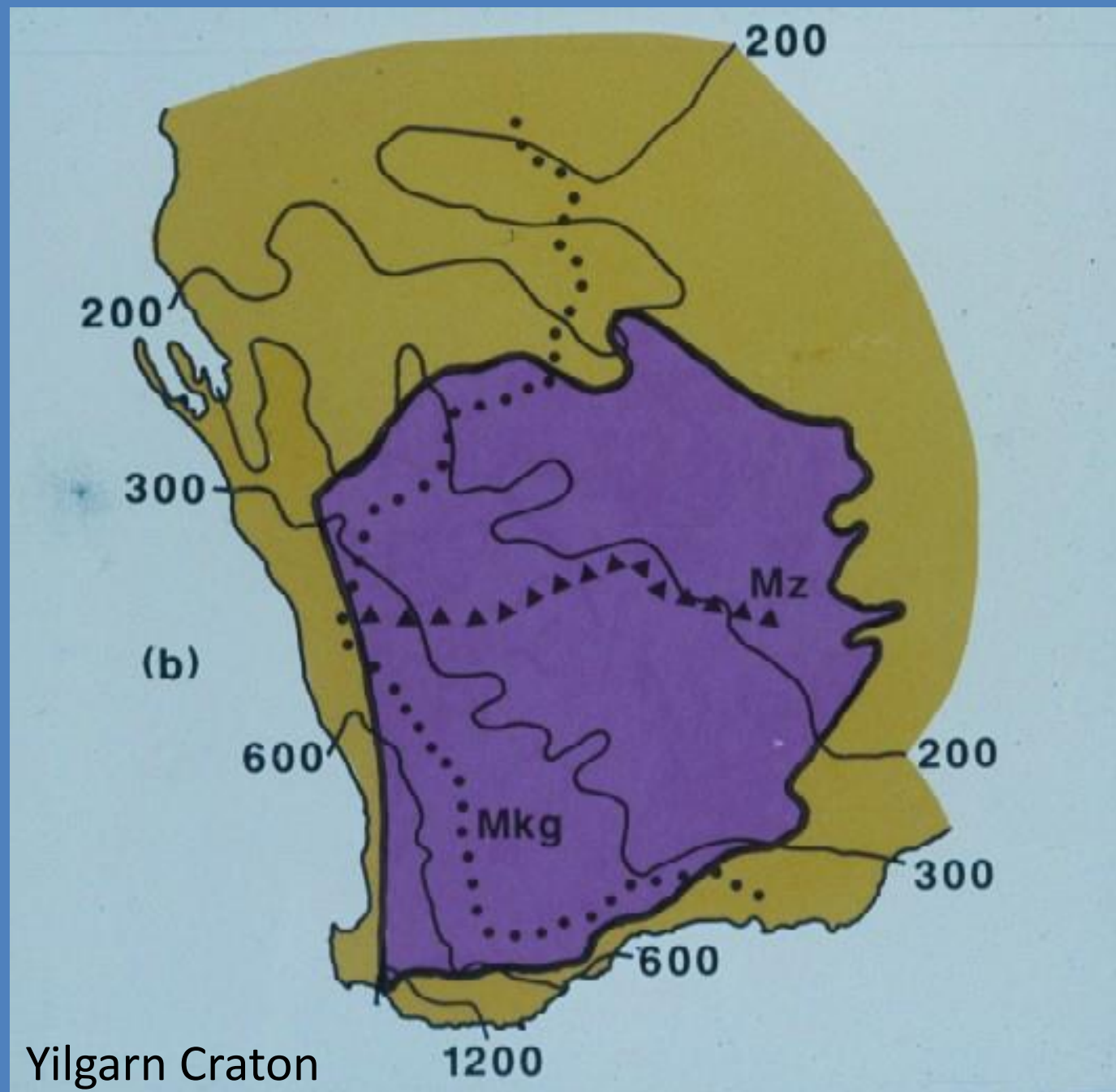
Time/ temperature/ water availability are interchangeable as factors in forming weathering. Tectonic stability is key in preserving weathering profiles.

Australia moved to lower latitudes before the Quaternary glaciations and escaped continental glaciation which, in many other continents, removed earlier weathering profiles.

However, some remnants of deep weathering are preserved, for example, in the Minnesota valley, Sweden, Finland, Scotland, Faeroe islands and are indicative of past weathering when the latitudinal temperature gradient was less than today.



(Yapp, 1987)



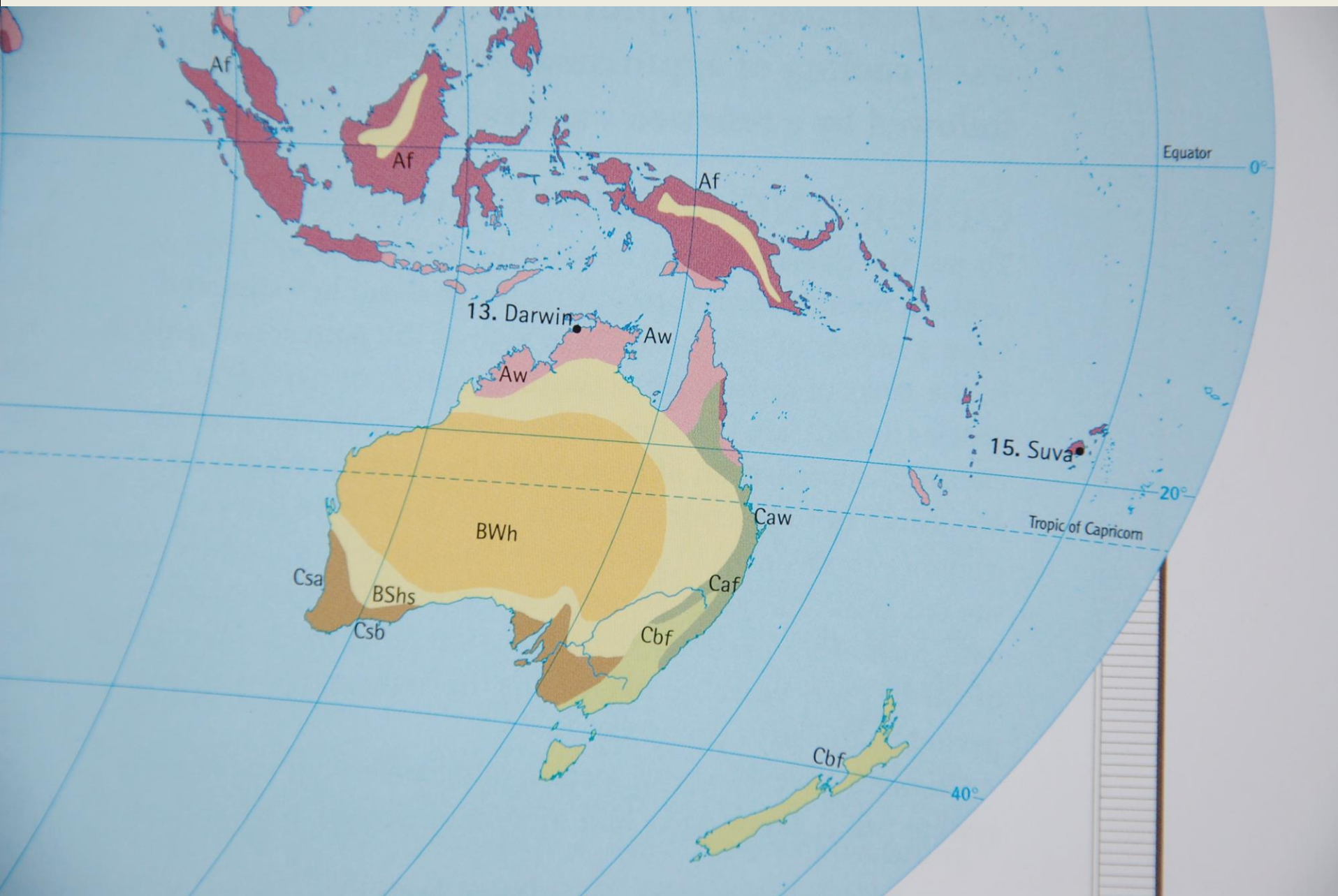
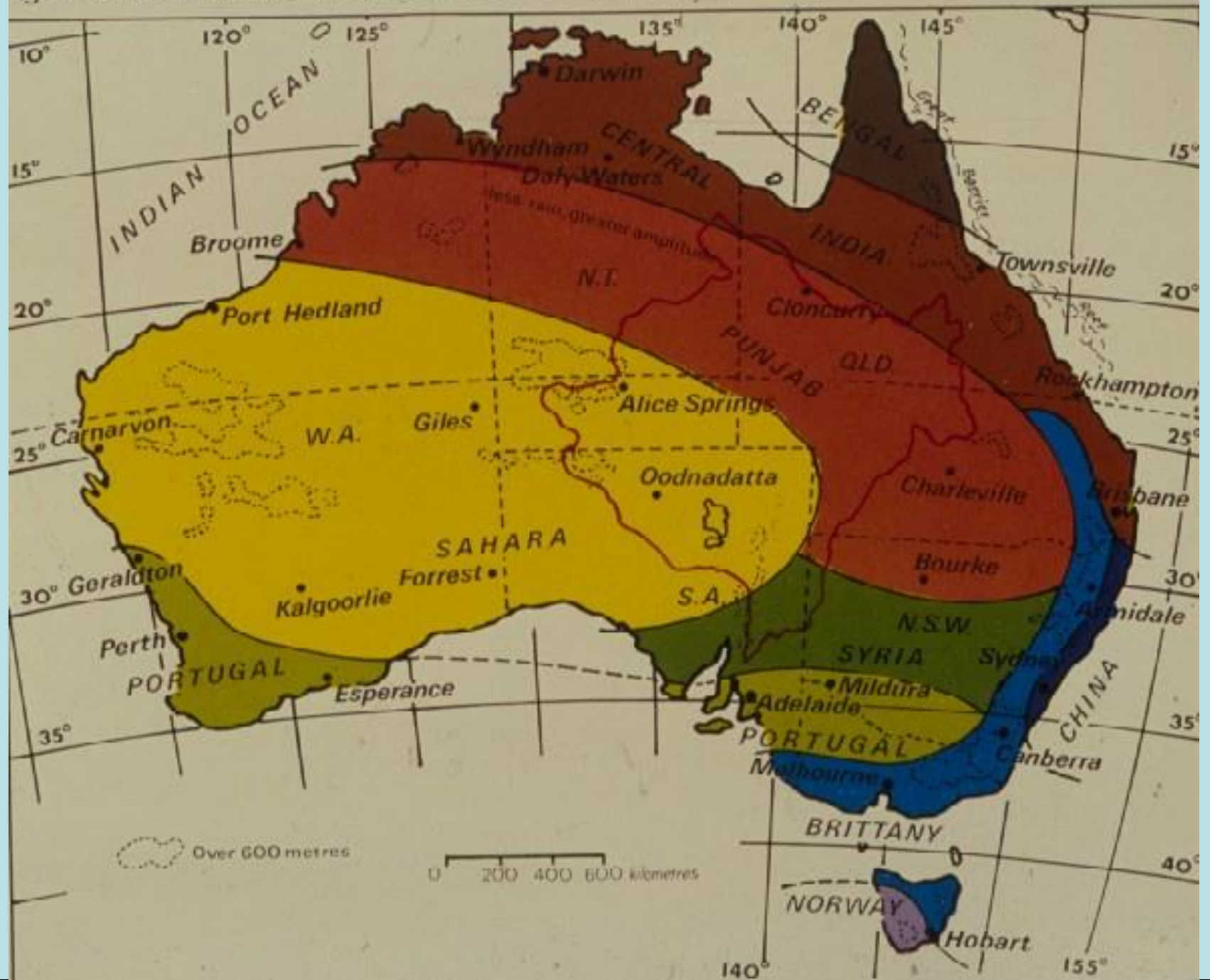
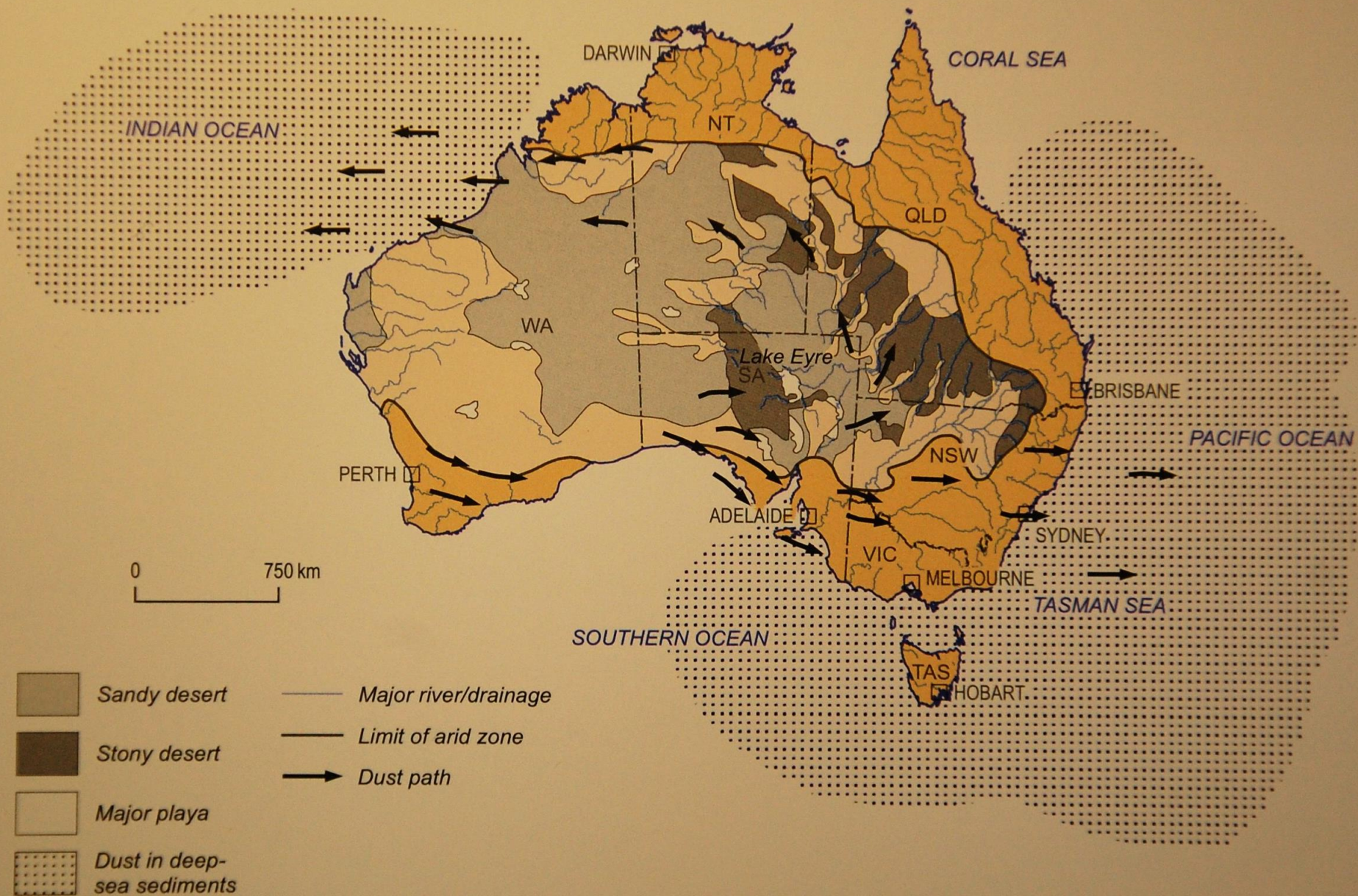


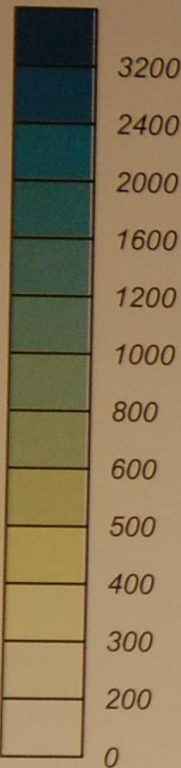
Figure 3. Climatic regions (after de Martonne, 1948)



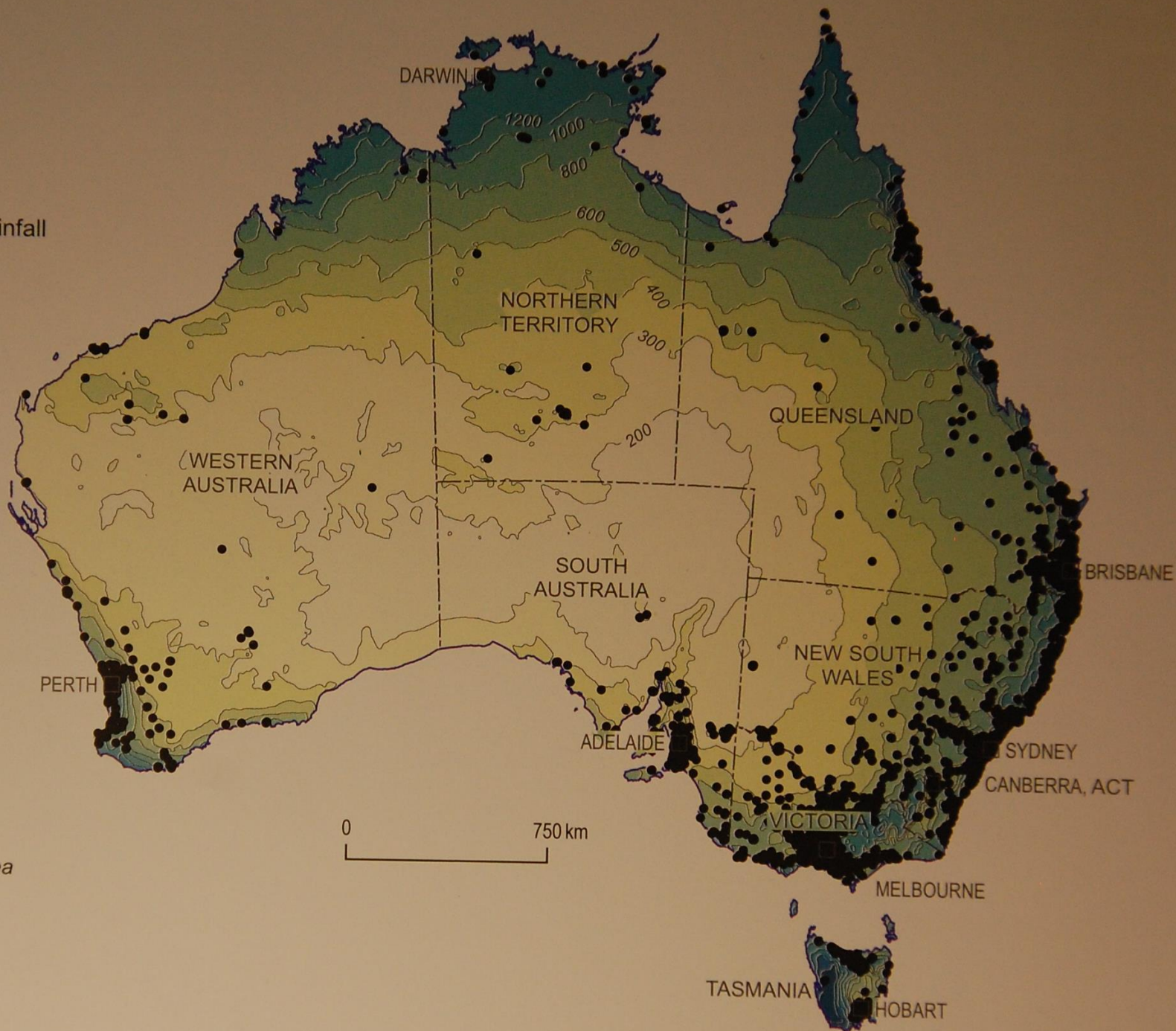




Average annual rainfall
(millimetres)



• *Populated area*

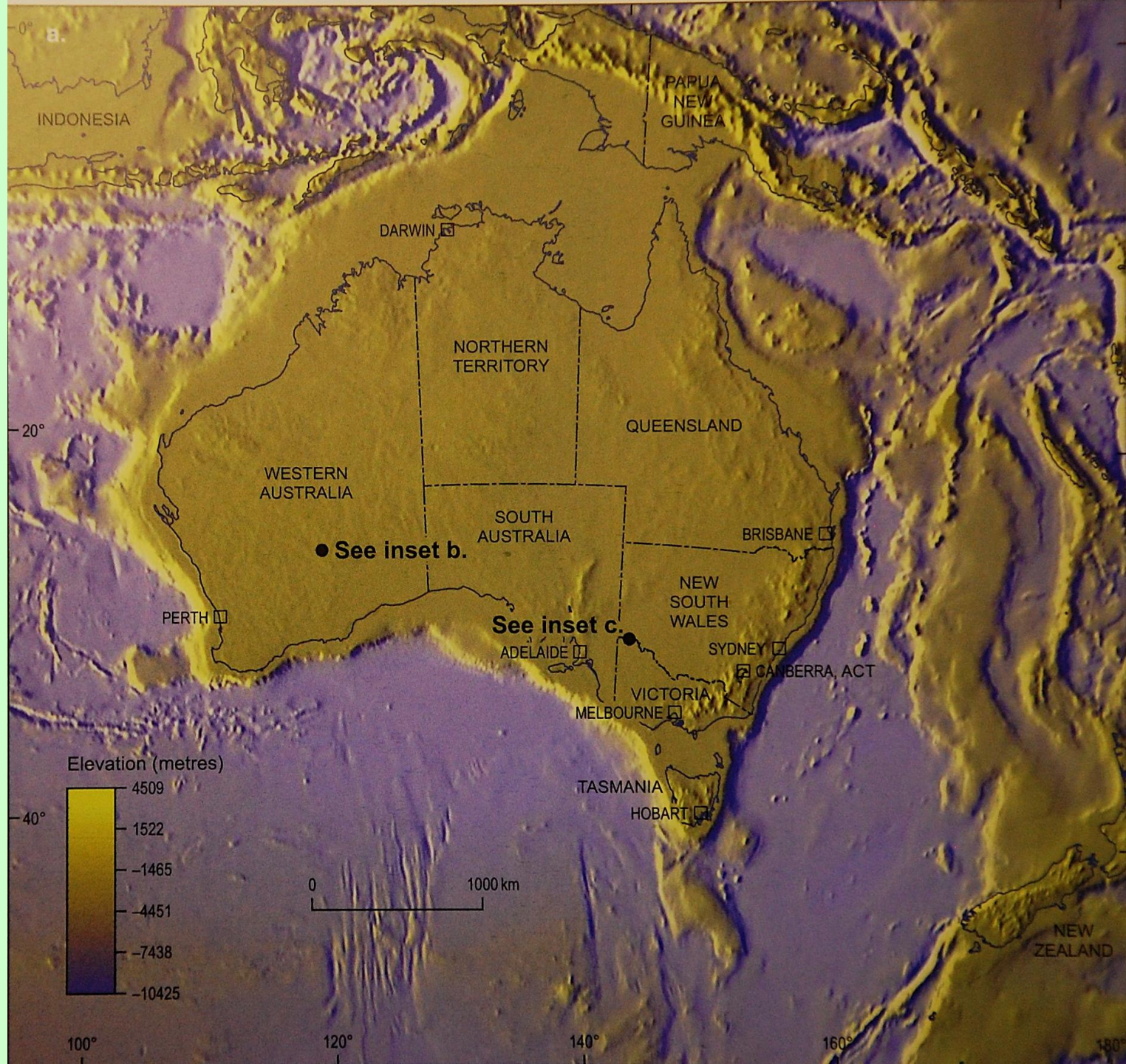


a.

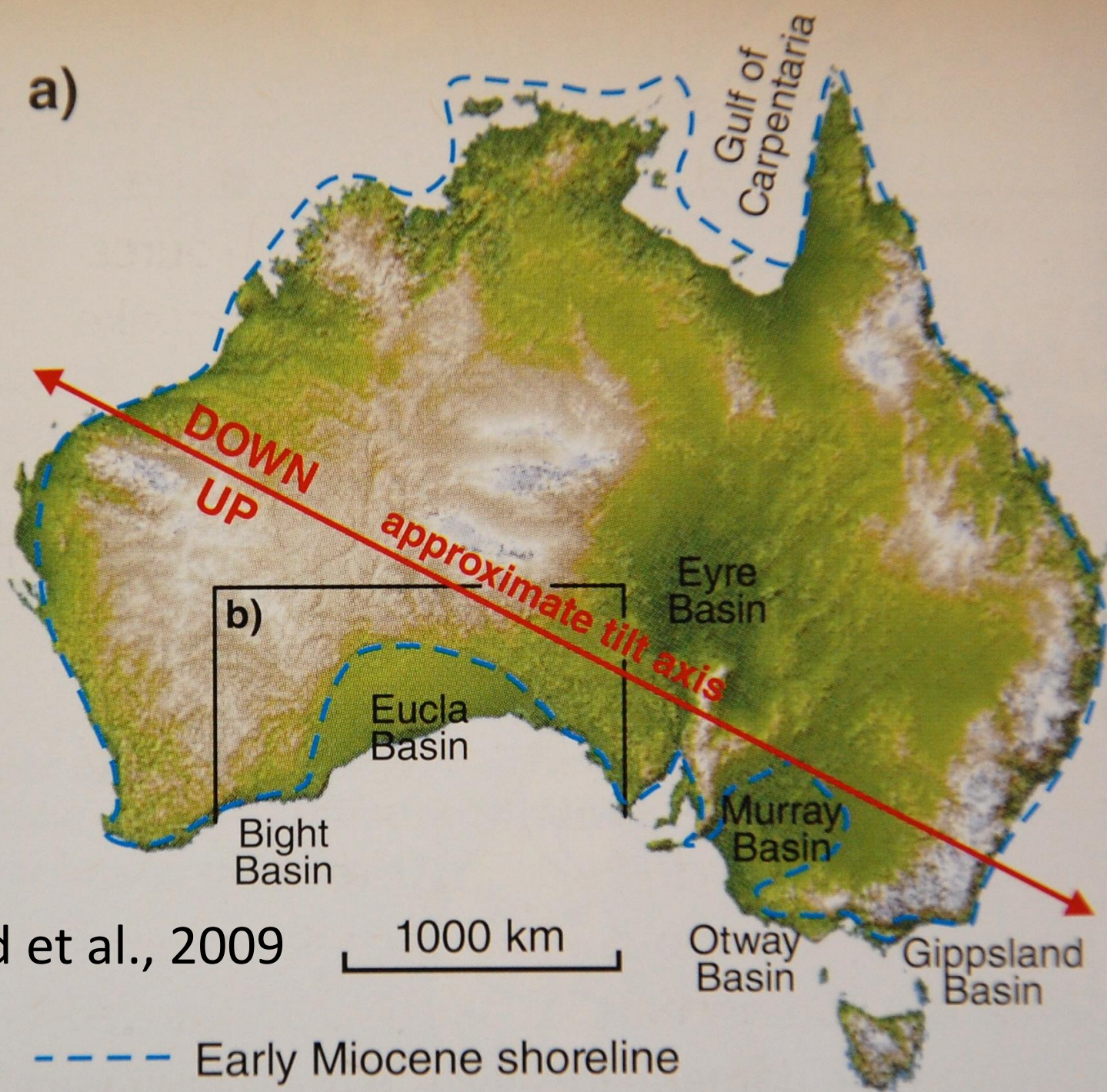








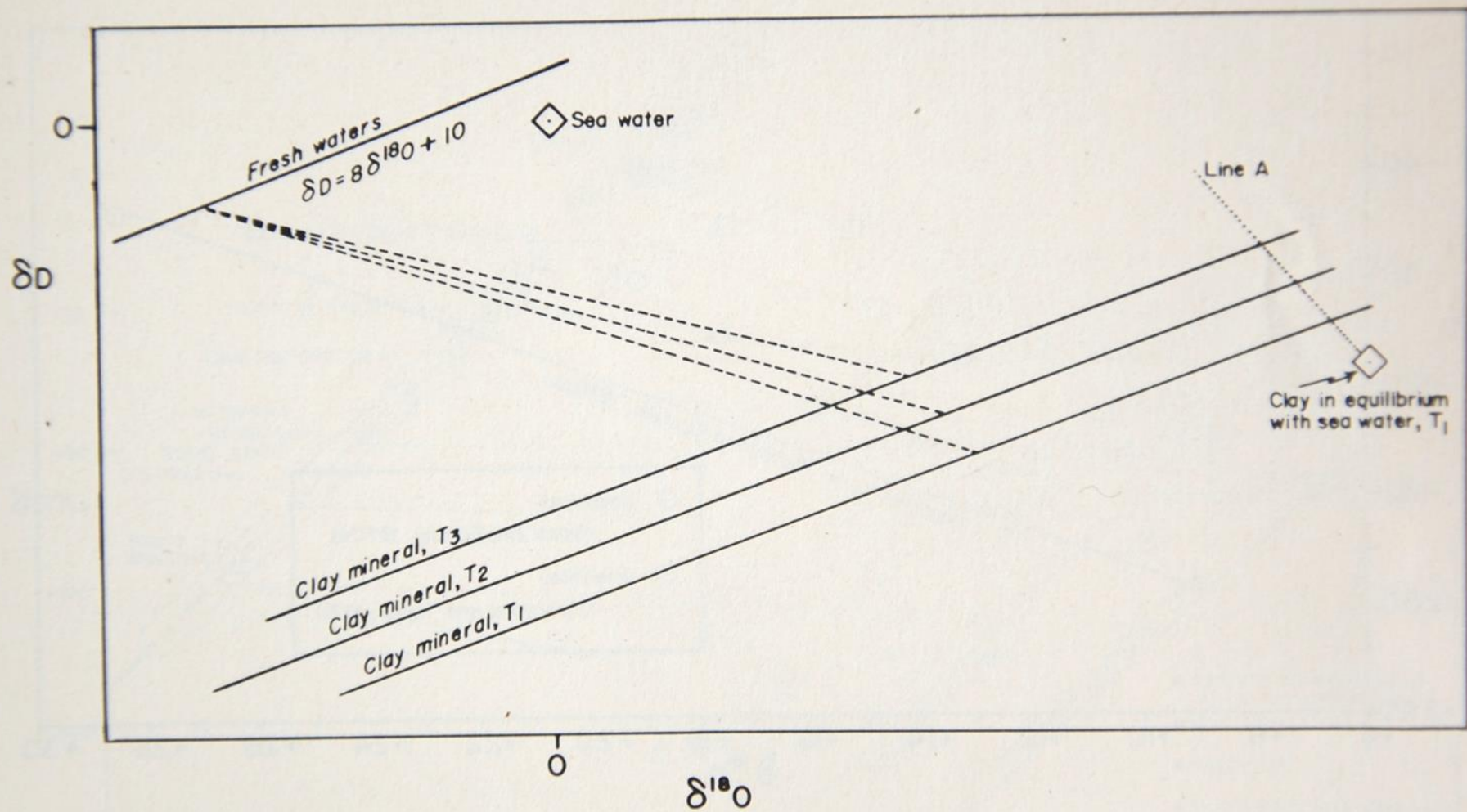
a)

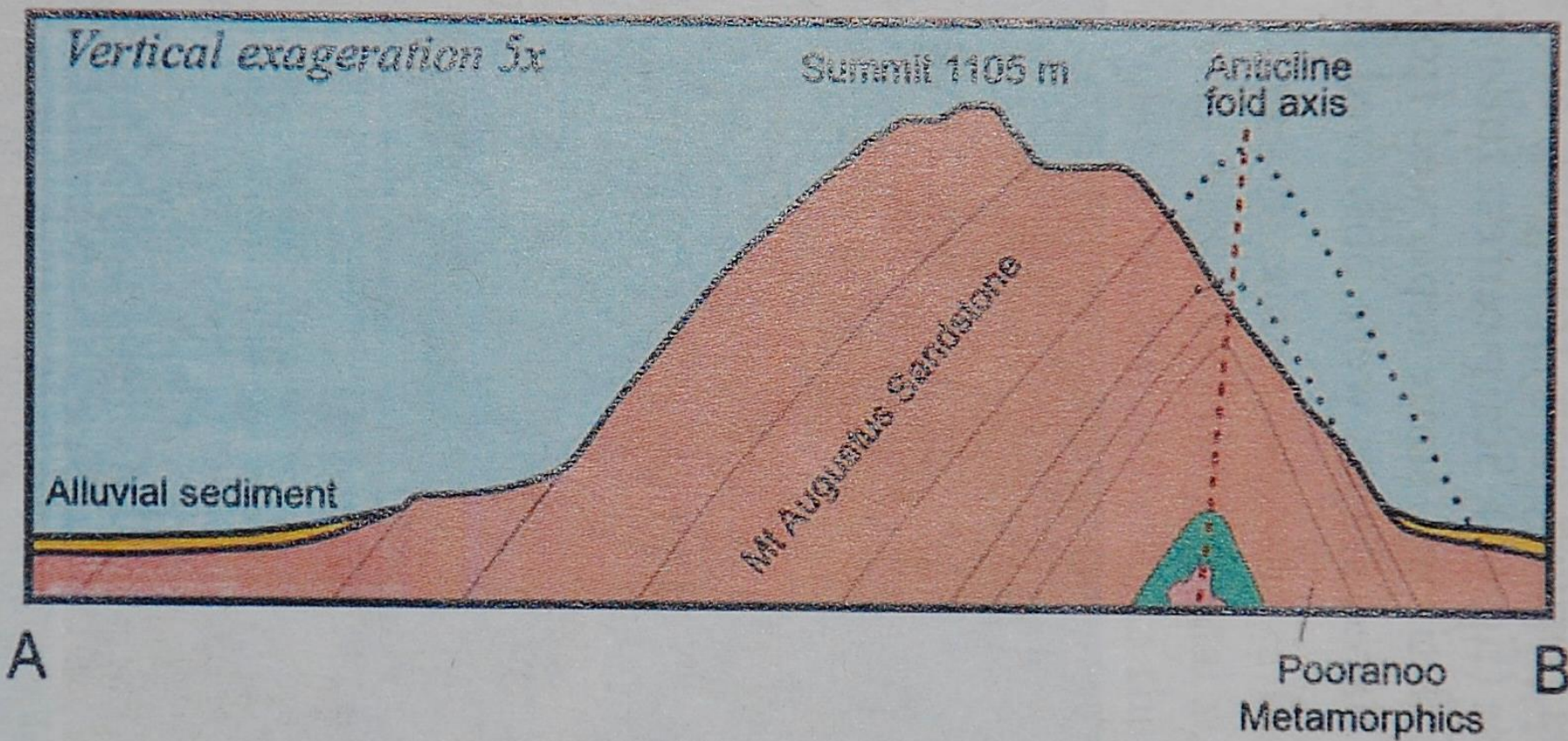


Sandiford et al., 2009











— Australian Plate boundary

- - - Plate boundary

... Plateau boundary

→ GPS track

→ Ridge push force

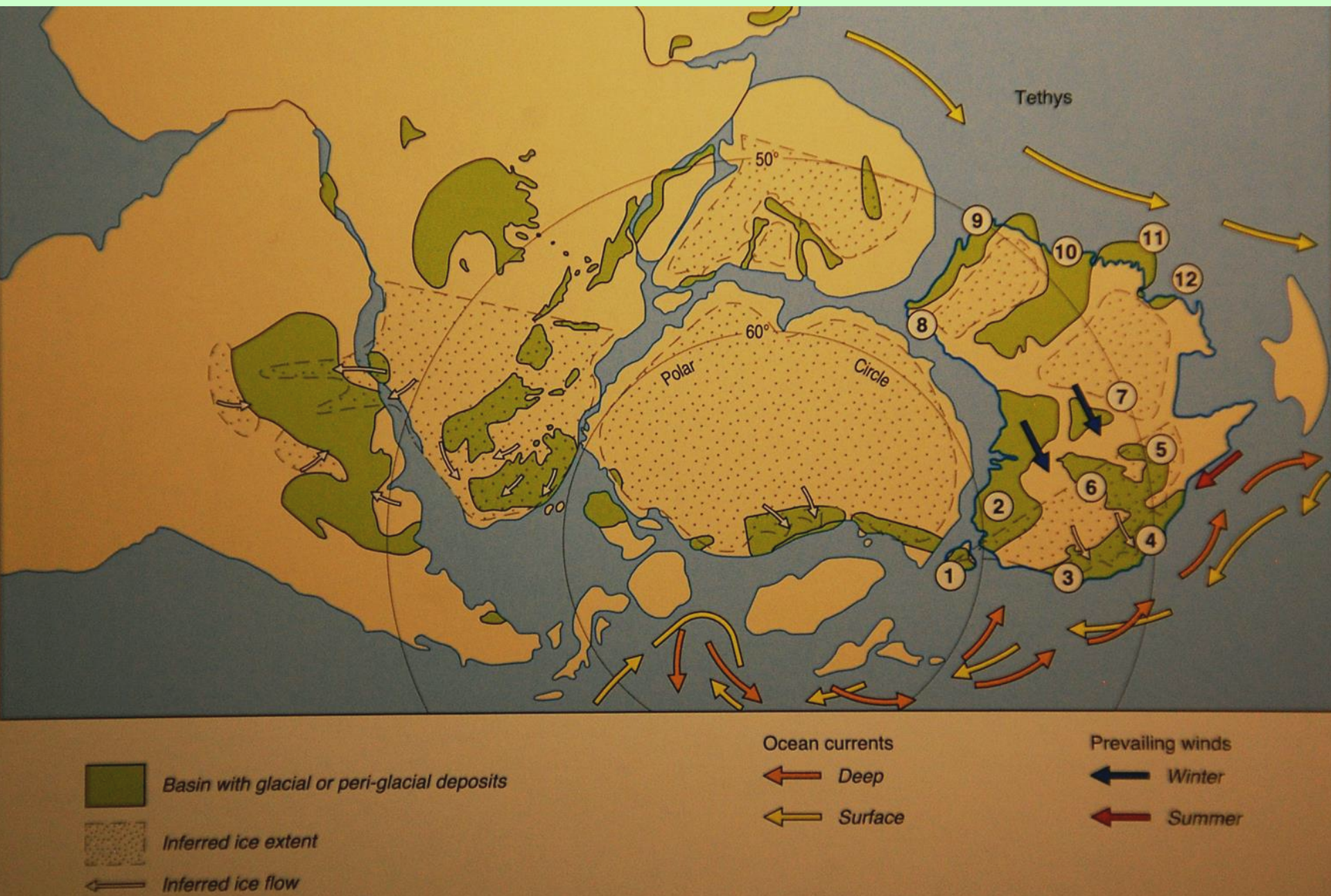
→ Plate driving force (subduction)

→ Plate resisting force (collision)

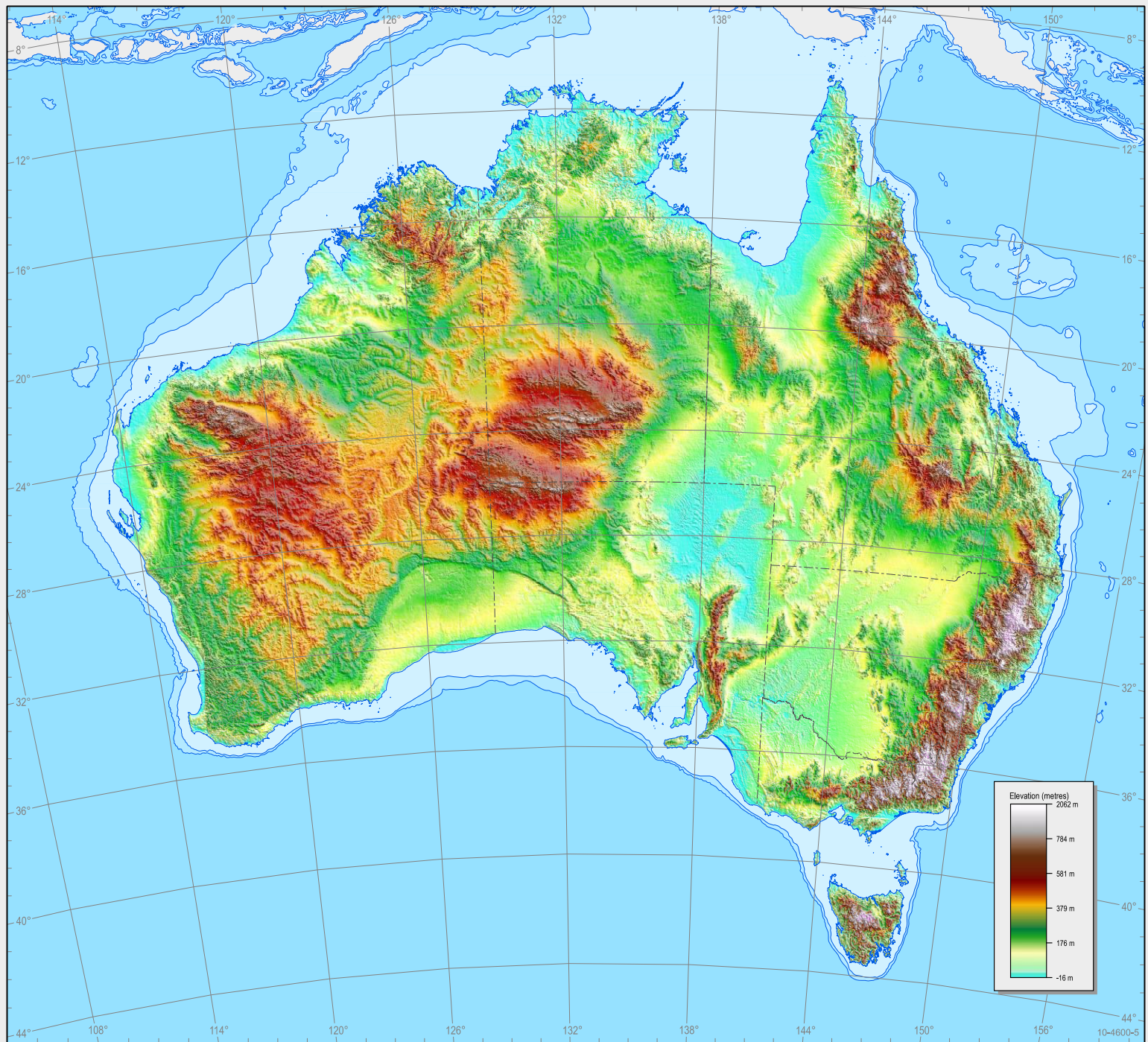
(ia) Island arc

(sz) Subduction zone

(cb) Collision boundary



Late Carboniferous – Early Permian (approx. 300 Ma)



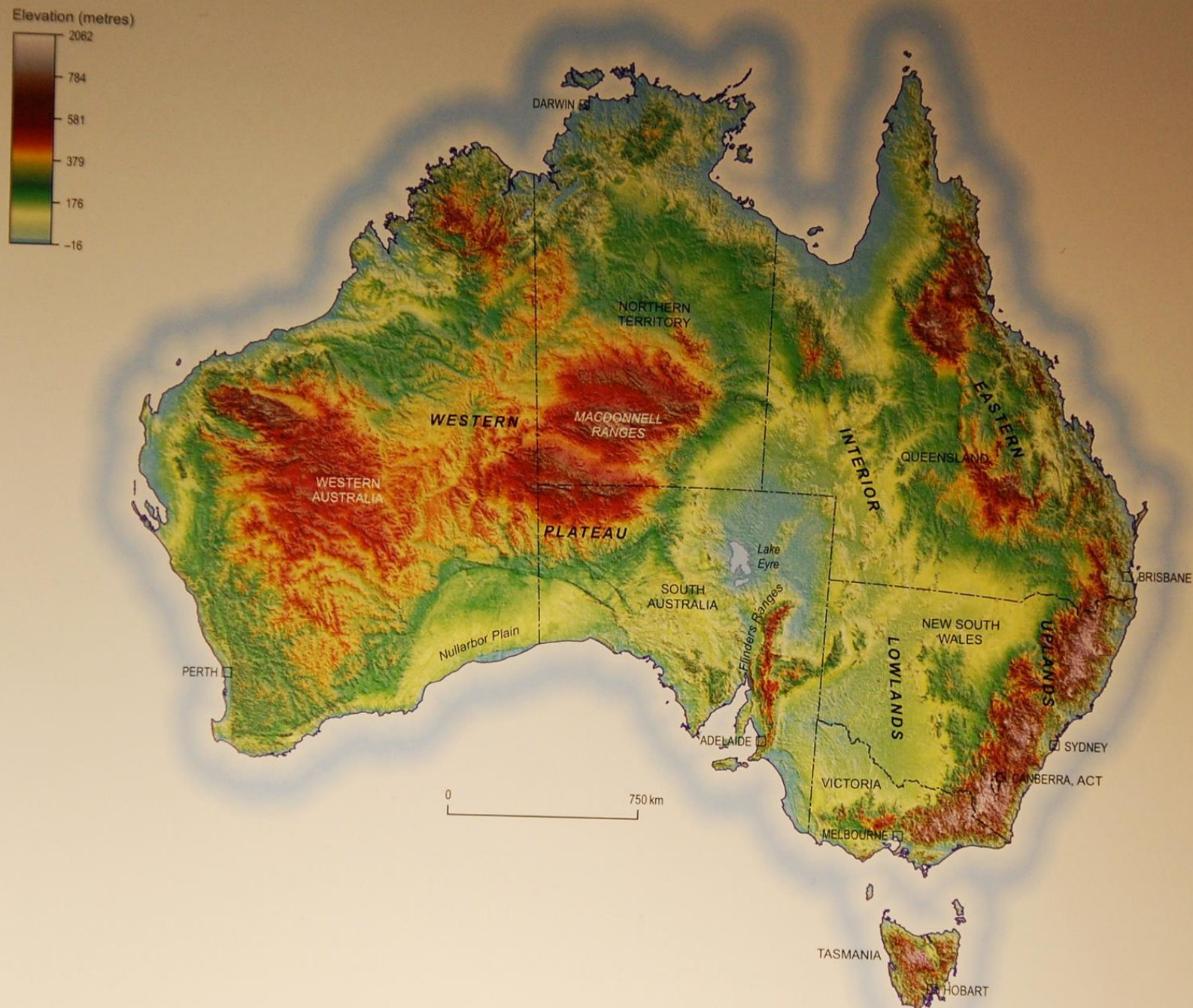
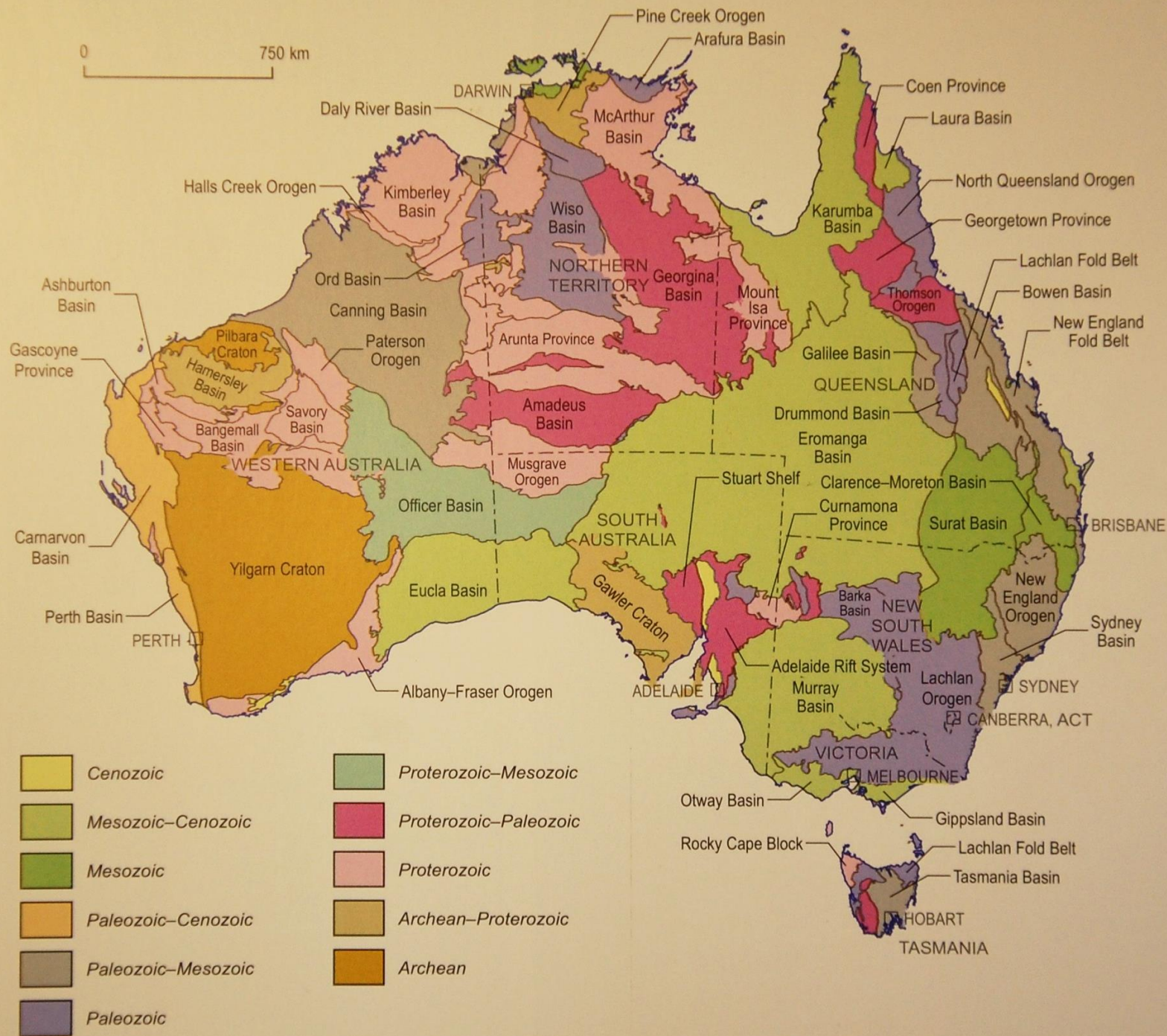
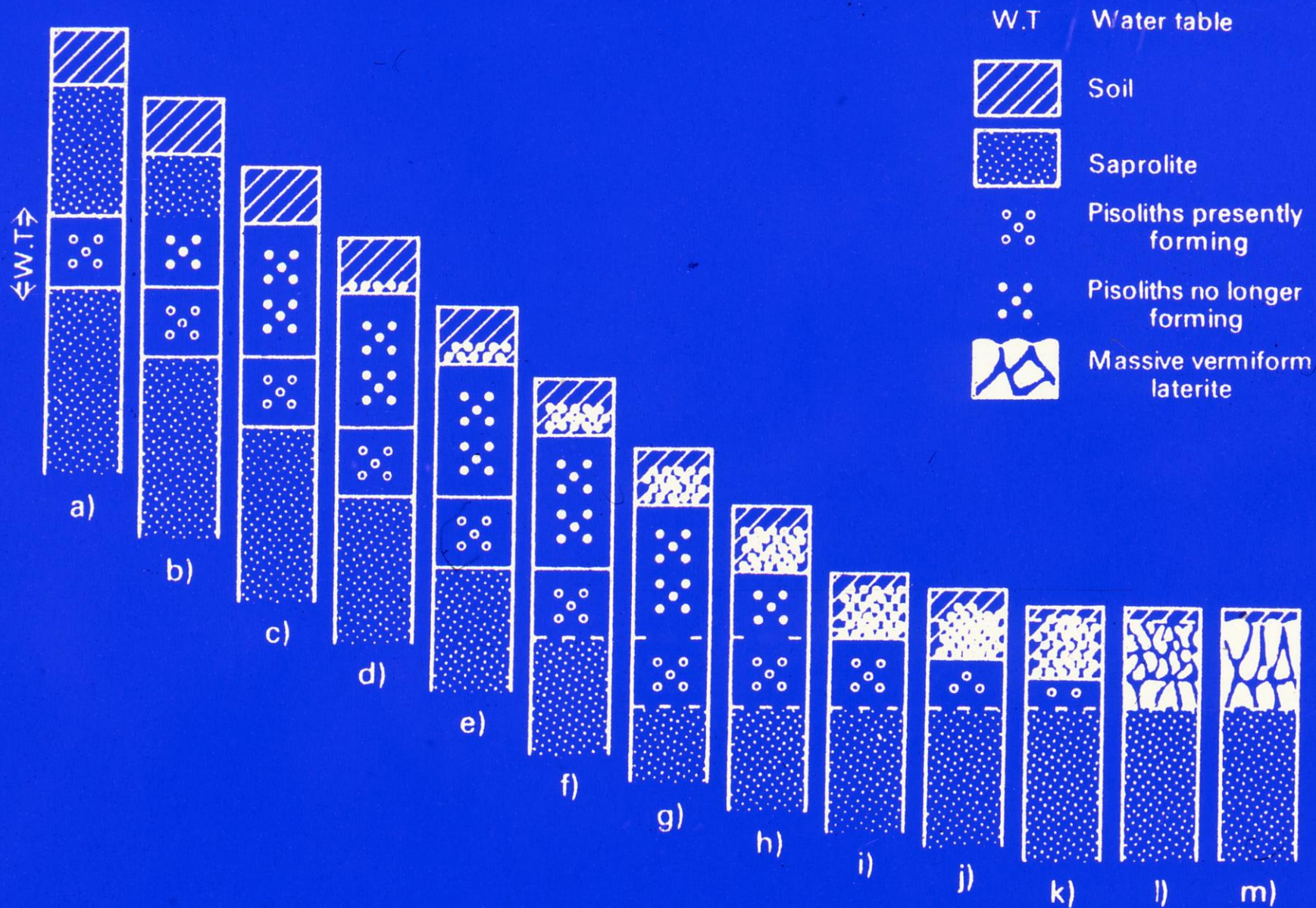


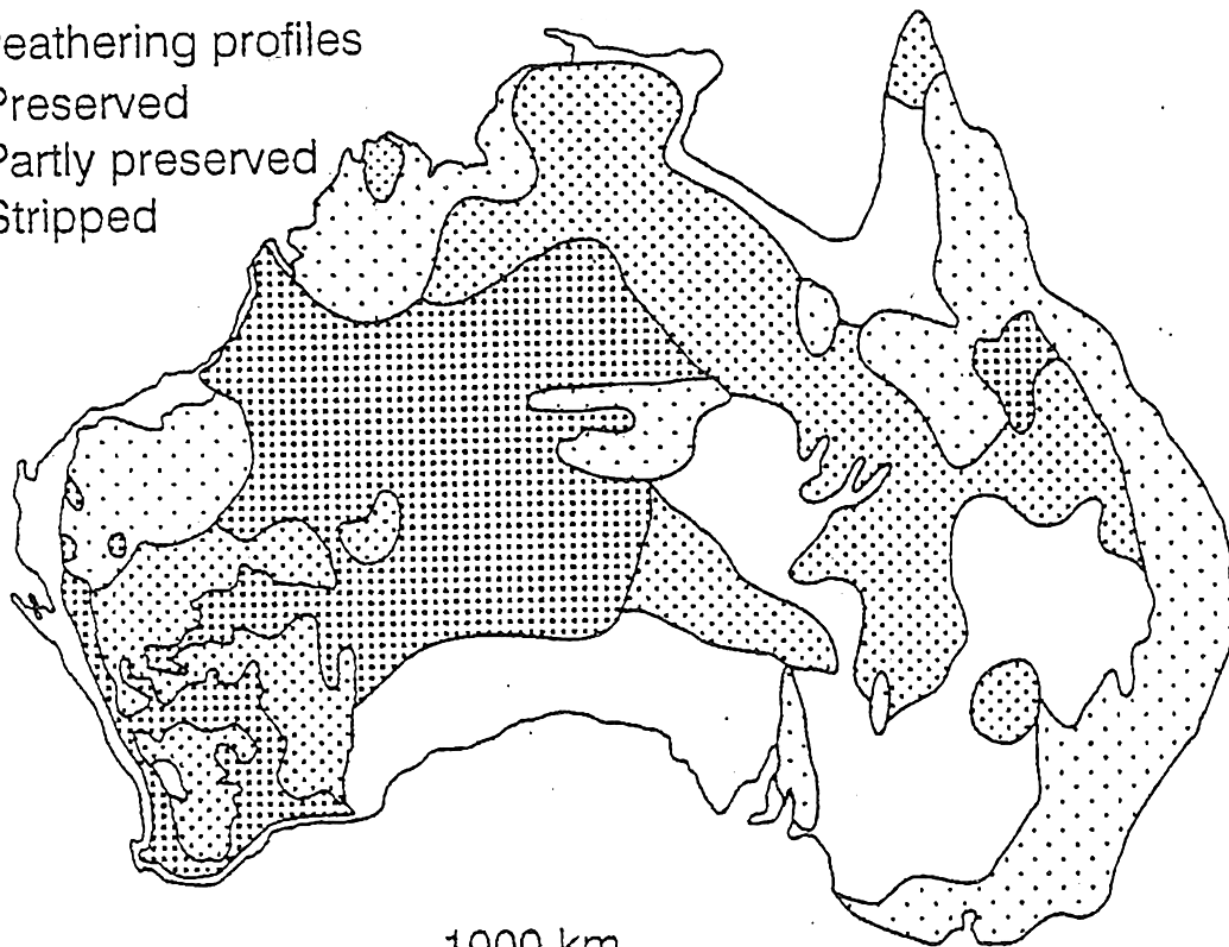
Figure 1.5: Digital elevation model of Australia, showing the main physiographic features of the continent.



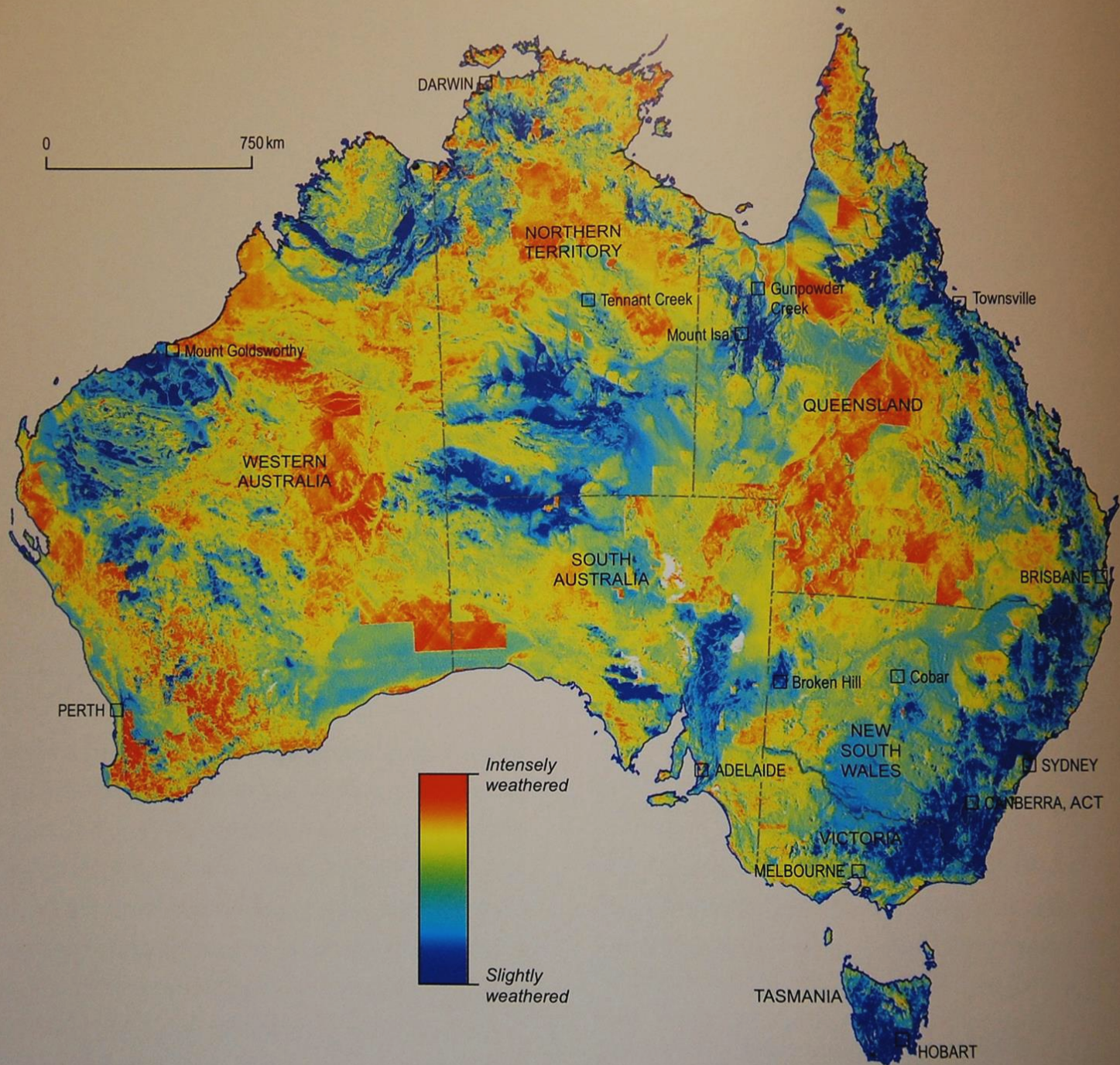


Deep weathering profiles

- Preserved
- Partly preserved
- Stripped



GTf217-00



Weipa, Queensland

